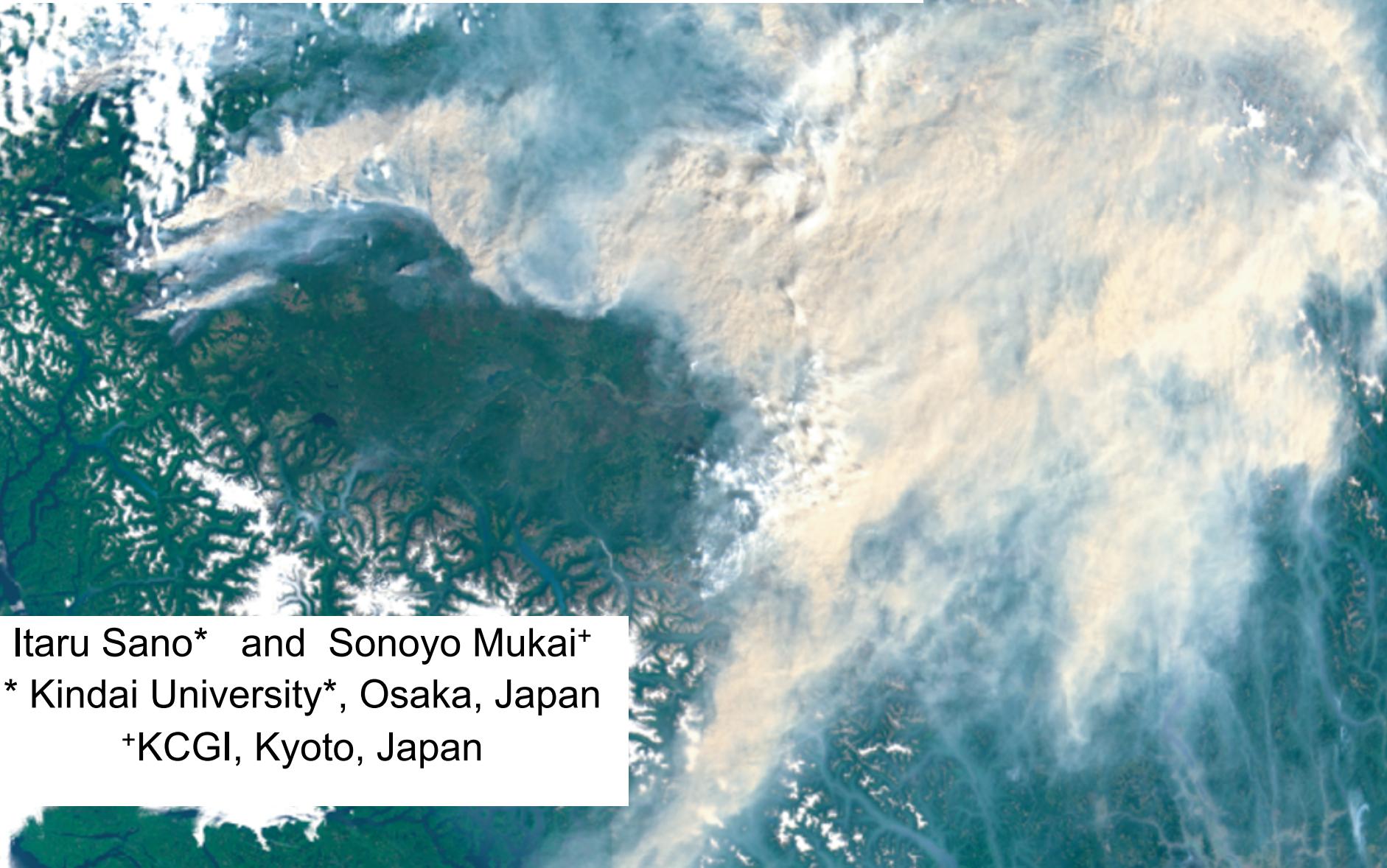


Detection of dense biomass burning area and the particle properties from GCOM-C / SGLI measurements



Itaru Sano* and Sonoyo Mukai[†]

* Kindai University*, Osaka, Japan

[†]KCGI, Kyoto, Japan

Background

Quantification of aerosol radiative forcing depends on the optical, chemical, and microphysical properties of carbonaceous aerosols.

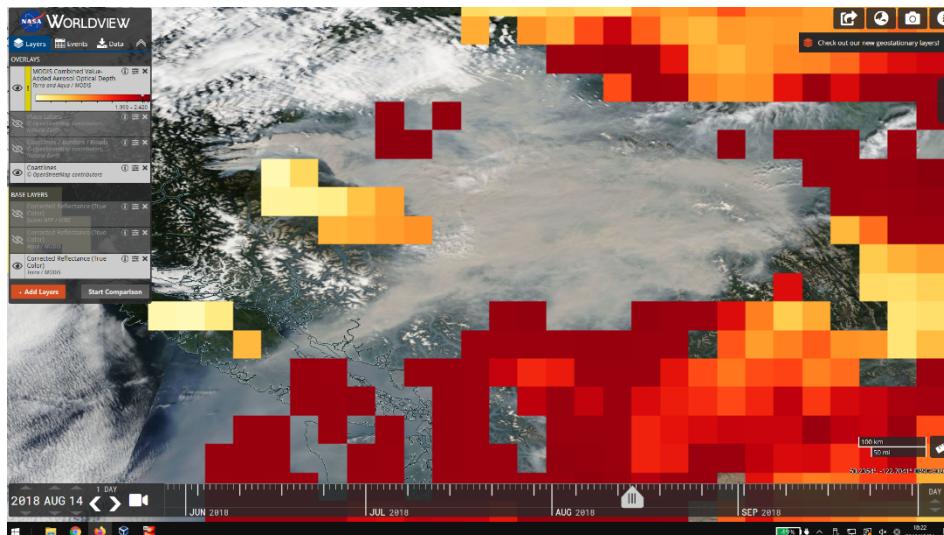
Nowadays, severe aerosol events, e.g., wildfire smokes, dust storms, and volcanic plumes, influence on ecosystem as health and air quality.



Worldview (MODIS) image on Aug. 14, 2018

Background

However, the details of such dense aerosols are not realized due to classify into haze or clouds.



Worldview (MODIS-AOD) image on Aug. 14, 2018

Objectives (#1)

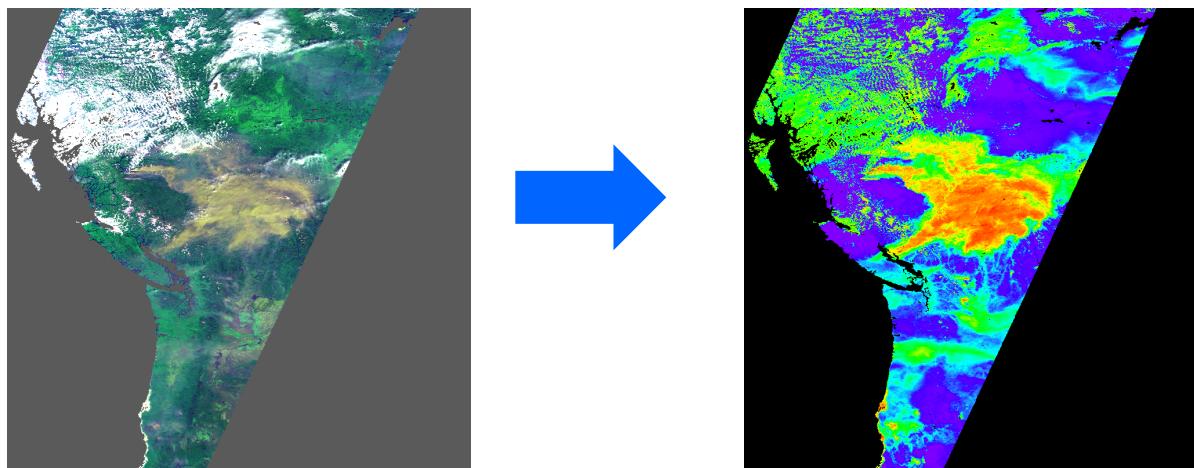
1. Detection and classification of absorbing aerosols

Biomass burning aerosols (BBA)

Mineral dusts (DUST)

Algorithm has been proposed based on

Japanese MODIS like sensor, GLI on ADEOS-2 (2002),
adapts to [Second GLI \(SGLI\) on GCOM-C \(2017\)](#).



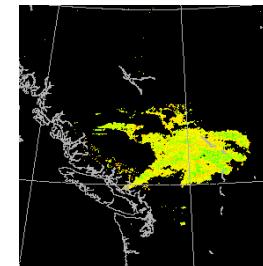
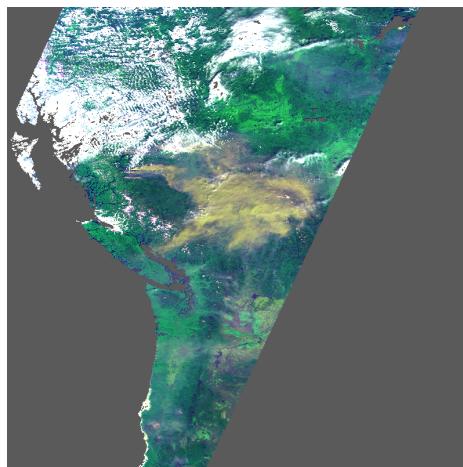
Objectives (#2)

2. Retrieving aerosol properties over dense aerosol region

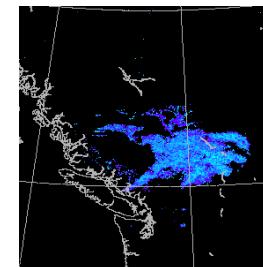
Ångström exponent (AE)

Single scattering albedo (SSA)

based on [semi-infinite atmosphere model](#) by
Vector version of Method of SOS (VMSOS).



AE



SSA

Classification of aerosol type

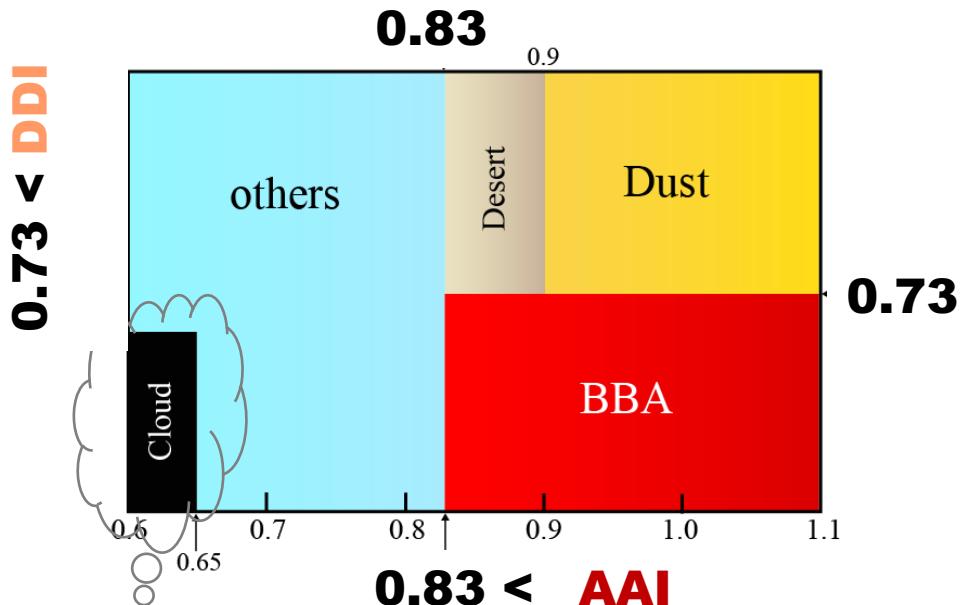
Absorbing aerosols in UV are discriminated from other type

$$\text{AAI} = R(412) / R(380) \quad \text{Absorbing aerosol index (AAI)}$$

Separation of dusts (Dust) from biomass burning aerosols (BBA)

$$\text{DDI} = R(2210) / R(380) \quad \text{Dust detection index (DDI)}$$

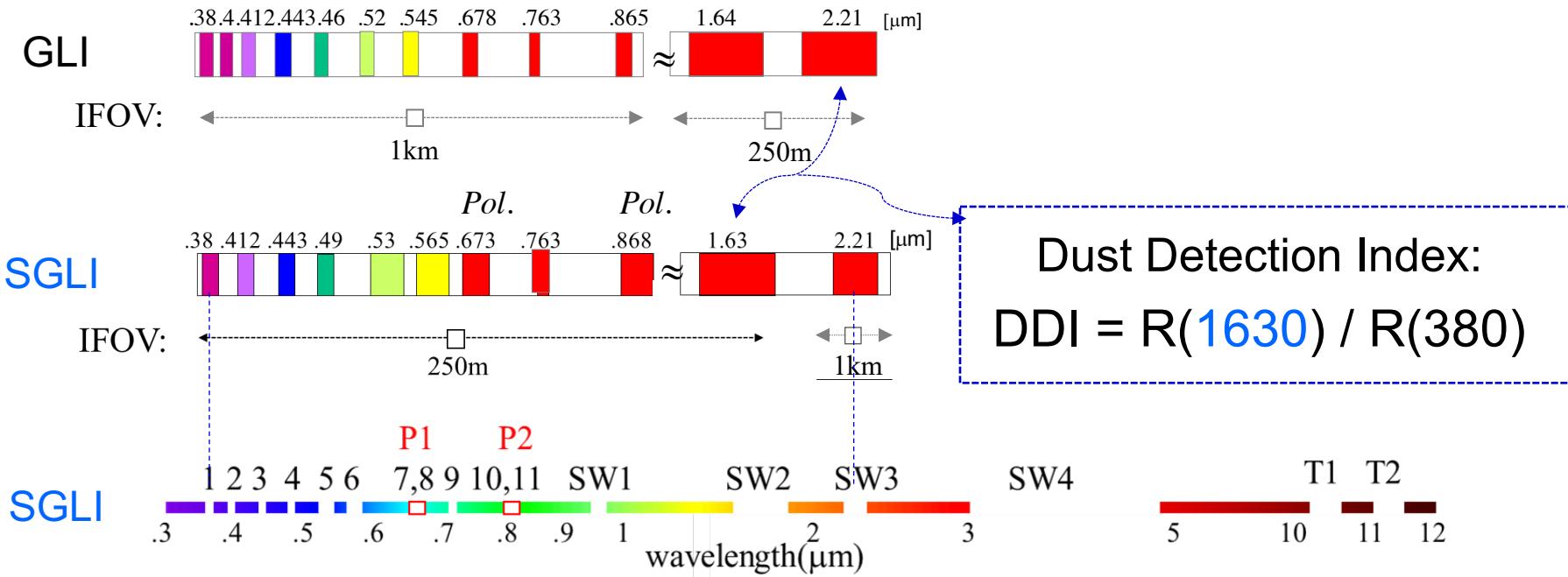
A discrimination chart
of aerosol types



Optically thick cloud
as COT(0.67) ≥ 20

* Mukai, S., Sano, I. and Nakata, M.:
J. Appl. Rem. Sen. 13(1), 014527,
doi:10.11117/1.JARS.13.014527, 2019.

Extension of aerosol type classification to SGLI / GCOM-C, 2017



In a similar manner to ADEOS-2/GLI (2002) case,
a discrimination chart of aerosol type classification
is examined using GCOM-C/SGLI data (2017).

Detection of absorbing aerosols : discrimination of BBA and DUST

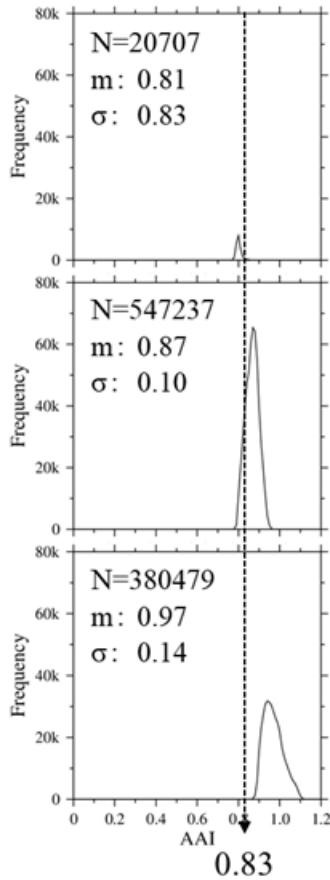
AOT < 0.1

0.1 ≤ AOT < 1.0

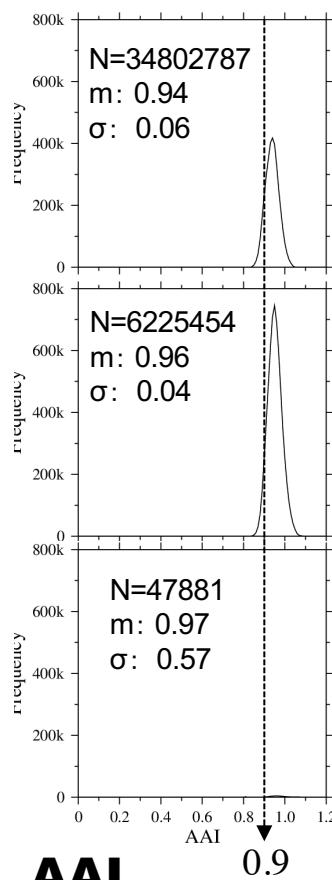
1.0 ≤ AOT

$$\text{AAI} = R(412) / R(380)$$

BBA

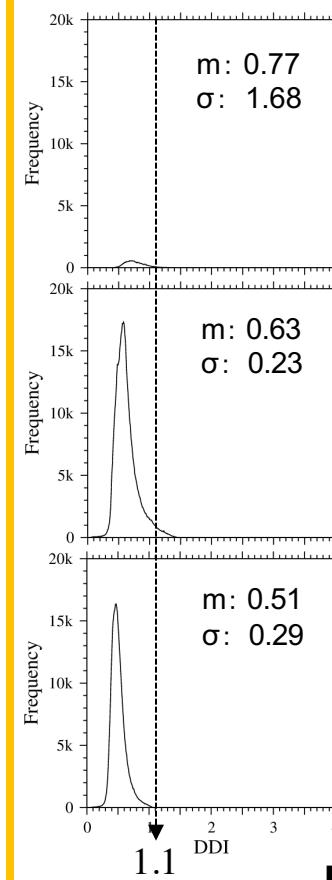


Dust

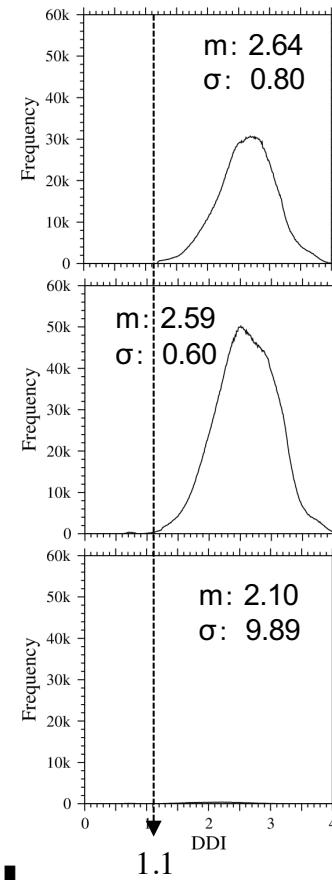


$$\text{DDI} = R(1630) / R(380)$$

BBA

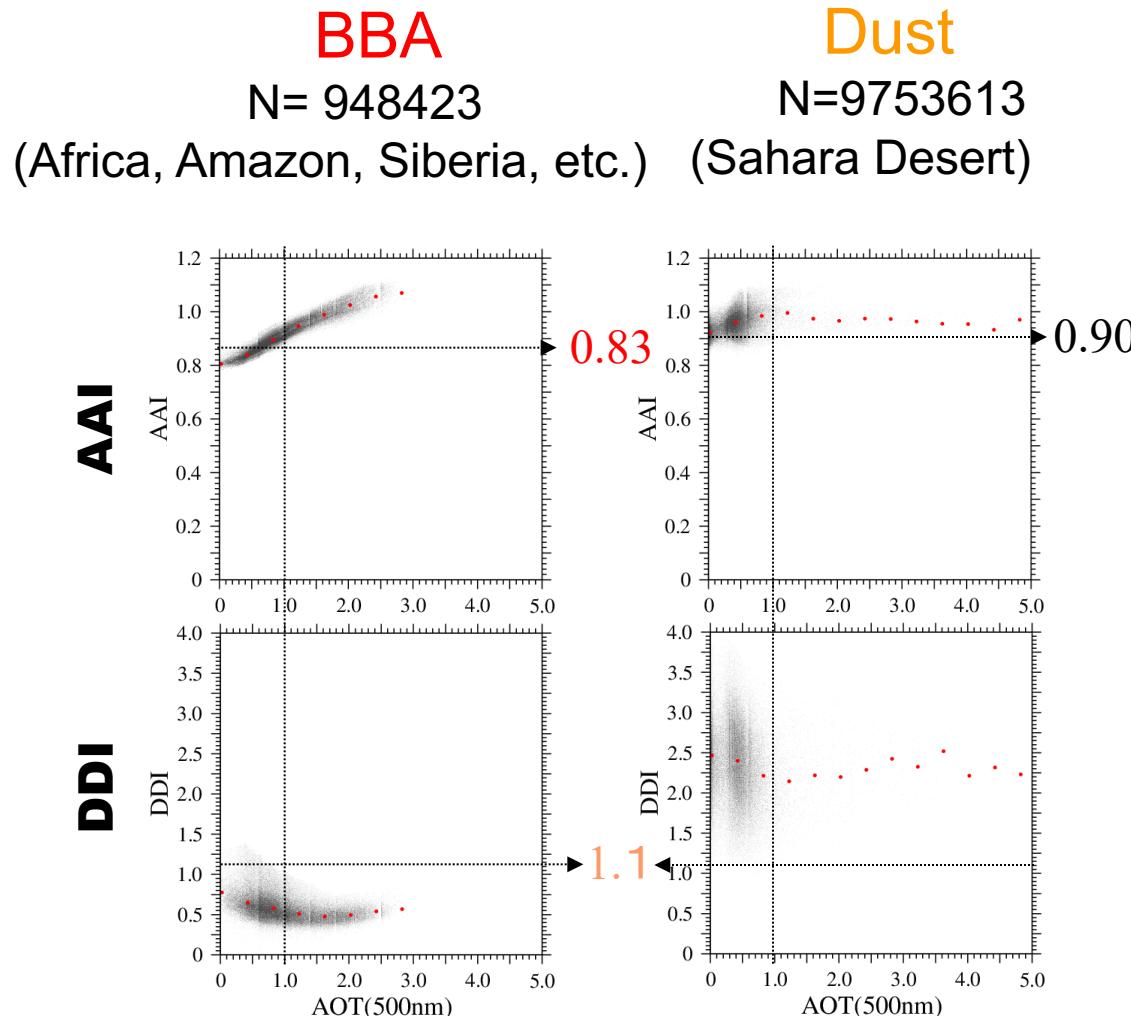


Dust



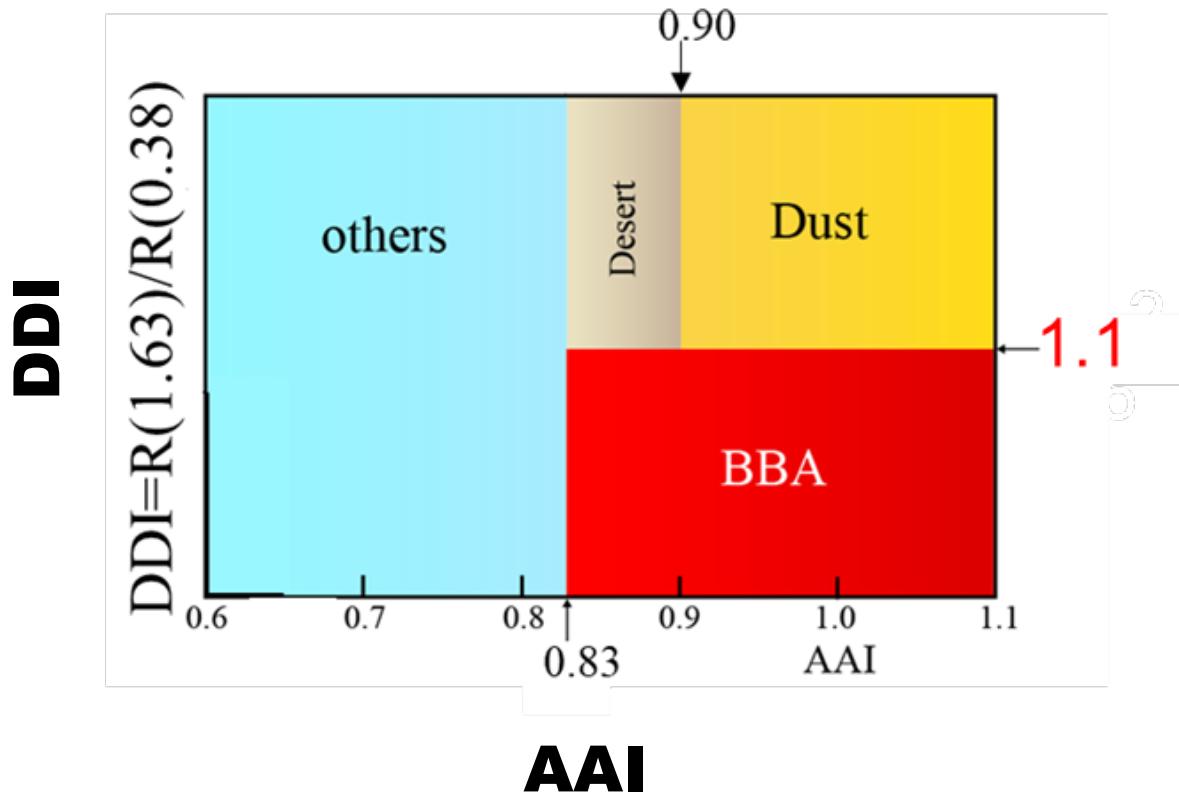
GCOM-C/SGLI (L2) data
from 1 May 2018 to 23 August 2019

Dependency of color ratio on AOT(500nm)

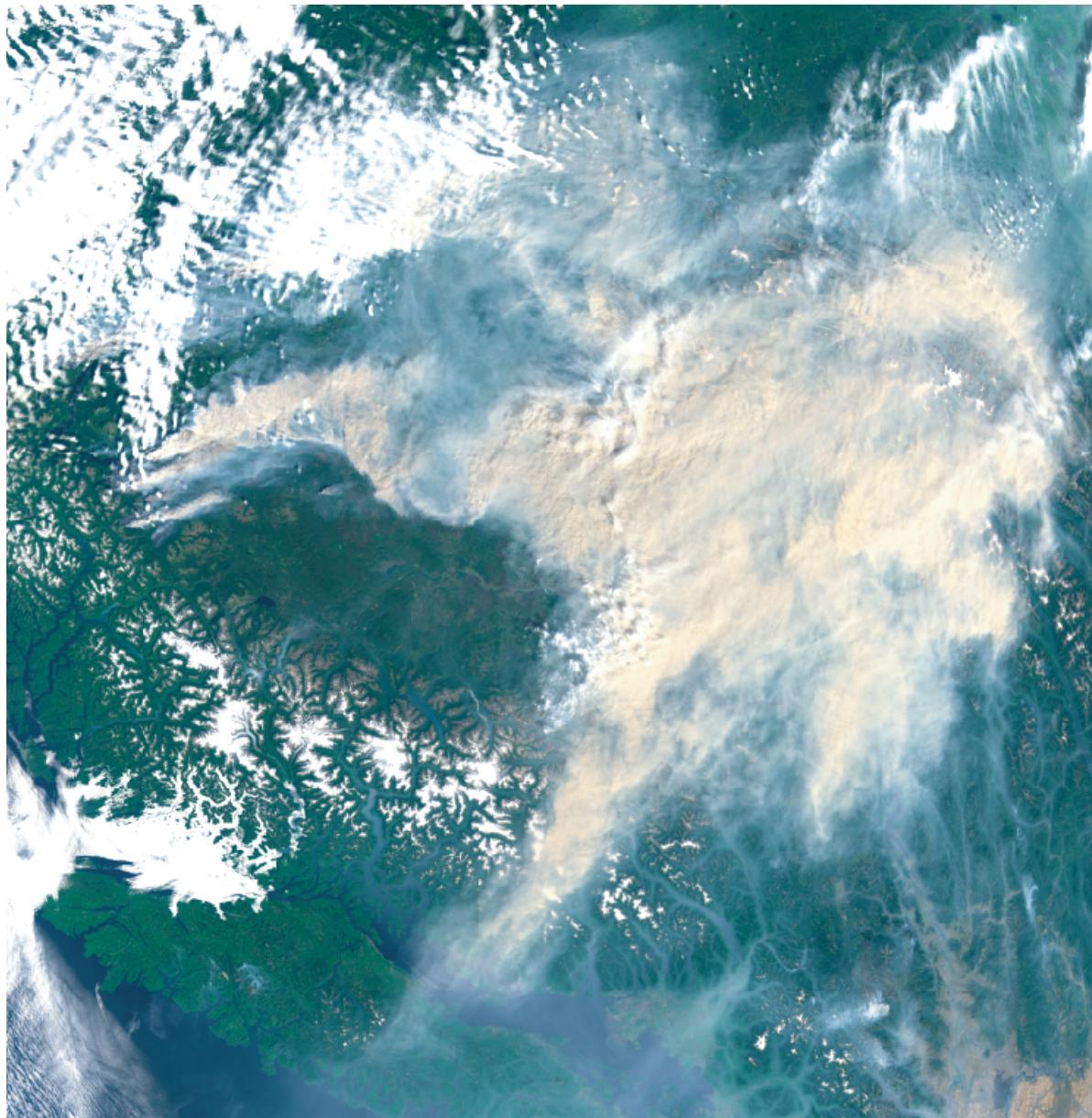


GCOM-C/SGLI (L2) data
from 1 May 2018 to 23 August 2019

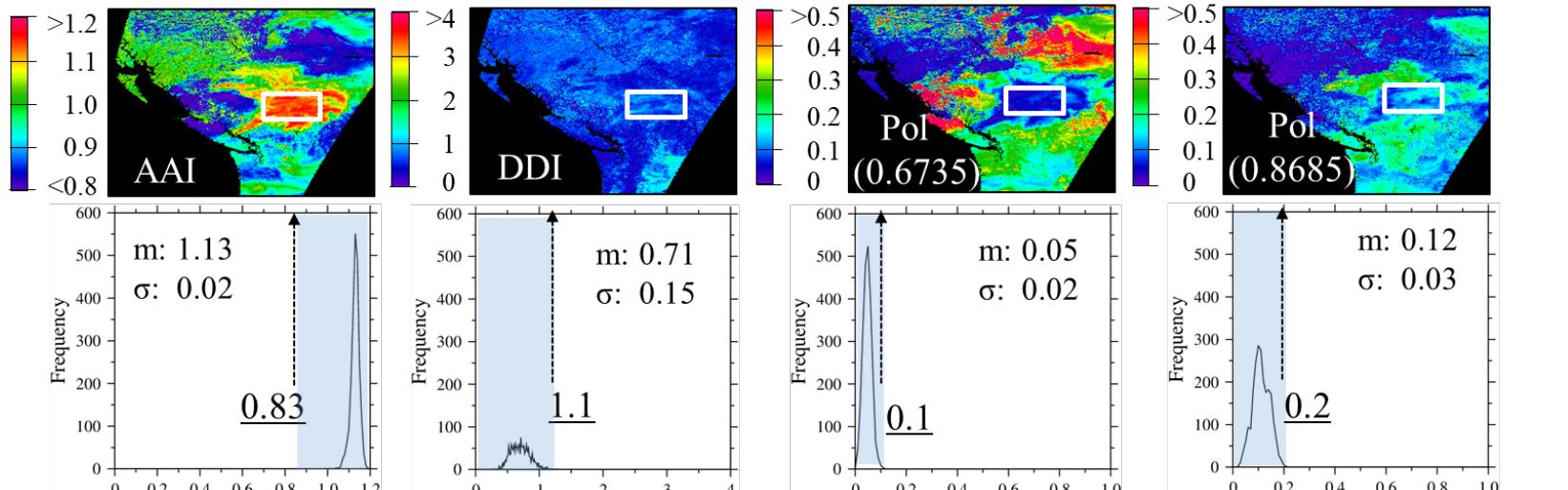
Discrimination chart of aerosols type from SGLI



SGLI imagery on Aug. 14, 2018



Characteristics of BBA type in Western Canada on Aug. 14, 2018



The BBA episode in western Canada coincides with the discrimination chart using AAI and DDI indices from SGLI global data.

#2 Aerosol retrieval part

Radiative Transfer in coupled atmosphere-surface system is roughly divided into 2-cases with respect to AOT(τ)

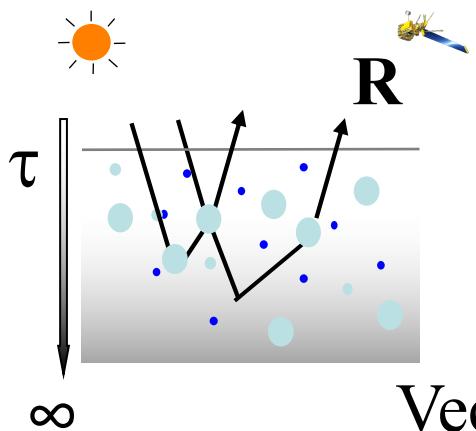
R. T. equation ;

$$\mu \frac{dI(\tau, \Omega)}{d\tau} = I(\tau, \Omega) - \frac{\omega}{4\pi} \int \tilde{P}(\Omega, \Omega') I(\tau, \Omega') d\Omega'$$

should be solved

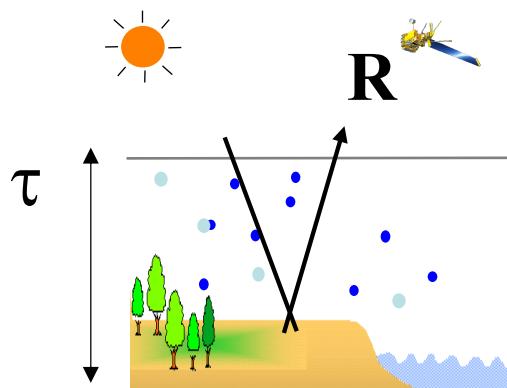
a) for infinite case: $\tau \sim \infty$

ex) aerosol events
dust storm /
biomass burning



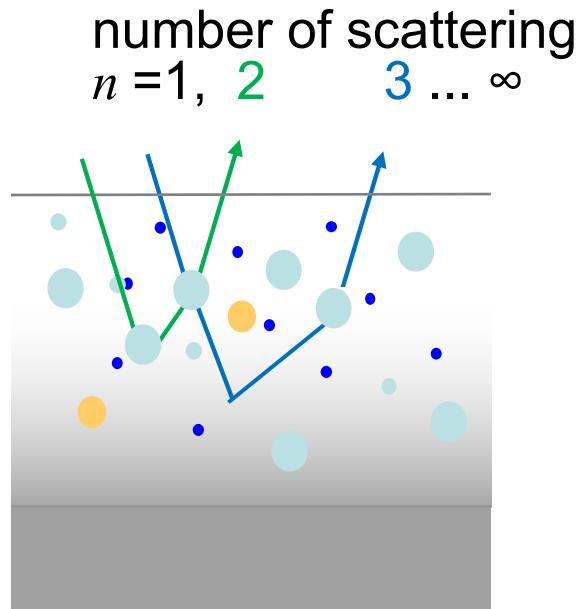
VMSOS:
Vector - Method of
Successive Order of Scattering

b) with boundary conditions for
finite atmosphere



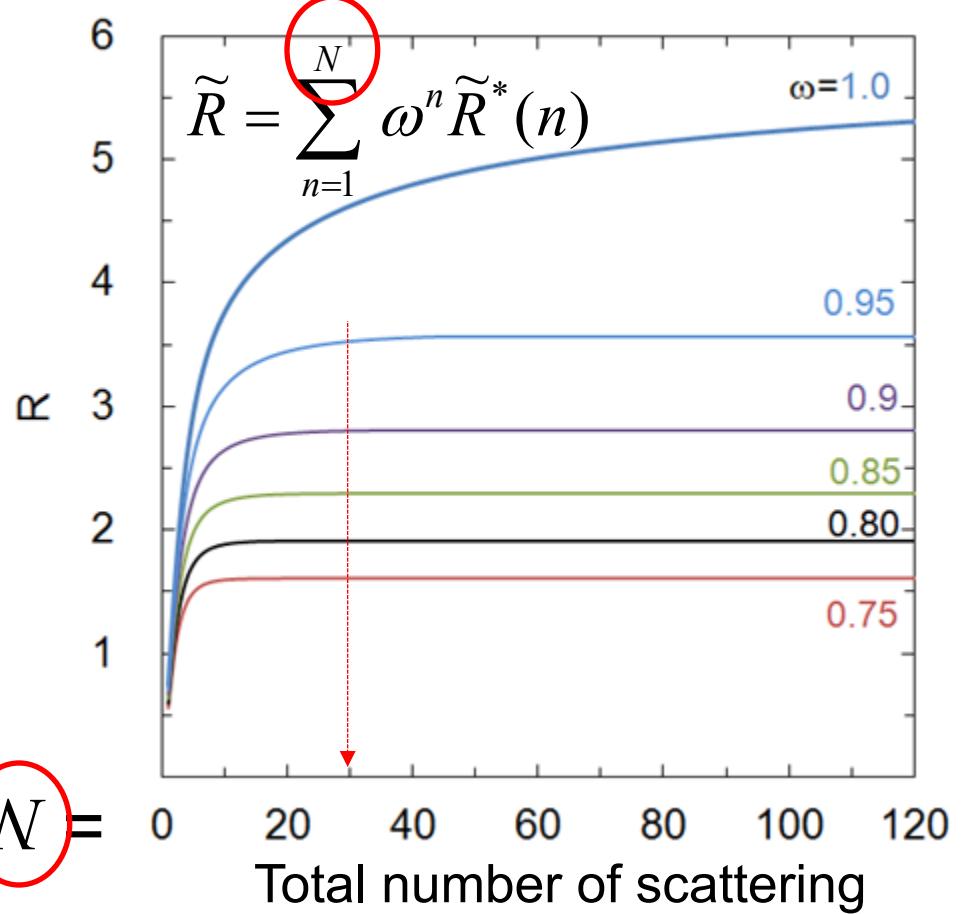
a) VMSOS; RT algorithm for semi-infinite atmosphere model in the polarization field

$$\tilde{R} = \sum_{i=1}^{\infty} \omega^n \tilde{R}^*(n)$$



VMSOS convergence
indication $\longrightarrow N$

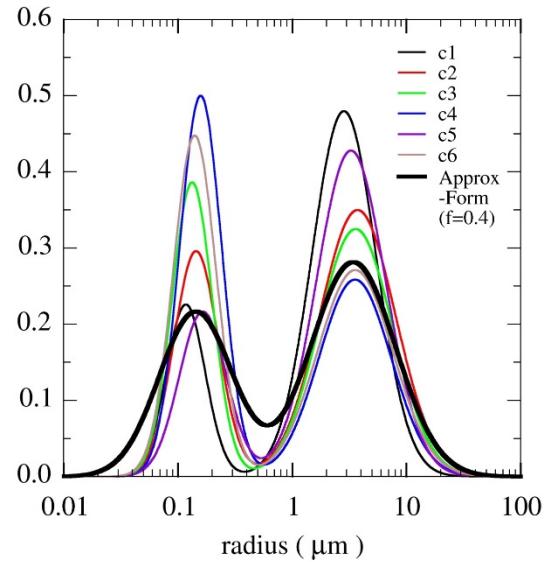
ω : albedo for single scattering
 $R^*(n)$: n^{th} - order of reflection fn.



Aerosol modeling based on AERONET

Size distribution : Bi - modal log-normal function

(Omar et al. JGR, 2005)



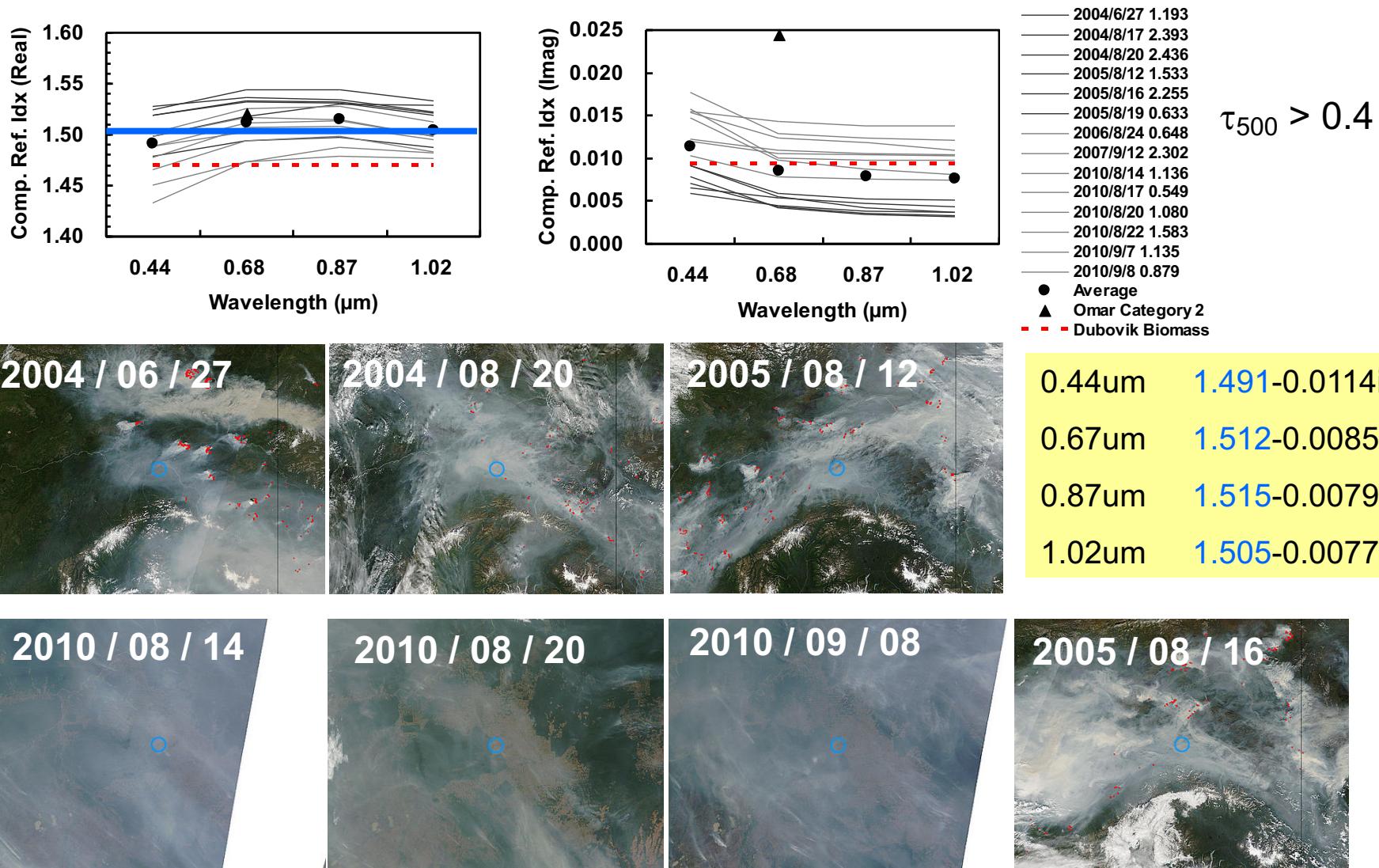
aerosol type	small		large	
	\bar{r}_f	σ_f	\bar{r}_c	σ_c
BB	0.144	1.562	3.733	2.144
RU	0.133	1.502	3.590	2.104
CP	0.158	1.526	3.547	2.065
DP	0.140	1.540	3.556	2.104
AF	0.144	1.533	3.607	2.104

Complex refractive index ($m = n - ki$) :

n : 1.50 (fixed)

k : 0.0001 - 0.1 (7 values)

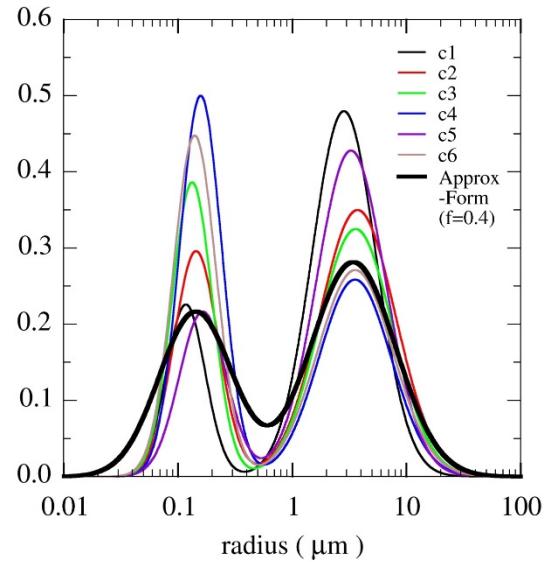
Optical properties of Biomass burning aerosols over Alaska (Bonanza), Amazon (JI_Parana_SE)



Aerosol modeling based on AERONET

Size distribution : Bi - modal log-normal function

(Omar et al. JGR, 2005)



aerosol type	small		large	
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Complex refractive index ($m = n - ki$)

n : 1.50 (fixed)

k : 0.0001 - 0.1 (7 values)

a) Aerosol retrieval by VMSOS : inputs and outputs

Inputs :

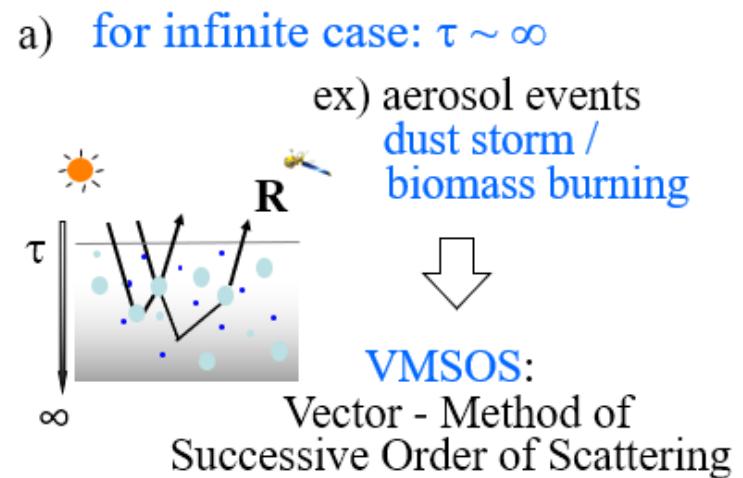
Nadir view SGLI measurements at 380, 412, and 530 nm.

Retrieved properties:

CMF of bimodal size dist.,
Imag. part of complex ref. idx.

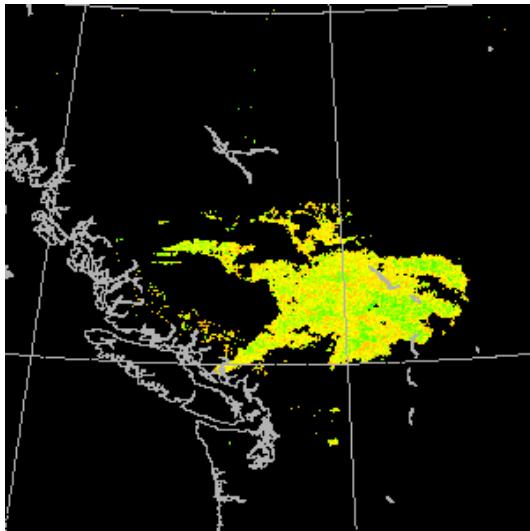
Outputs:

→ **Angstrom Exponent (AE),**
and **SSA (380, 530 nm)**

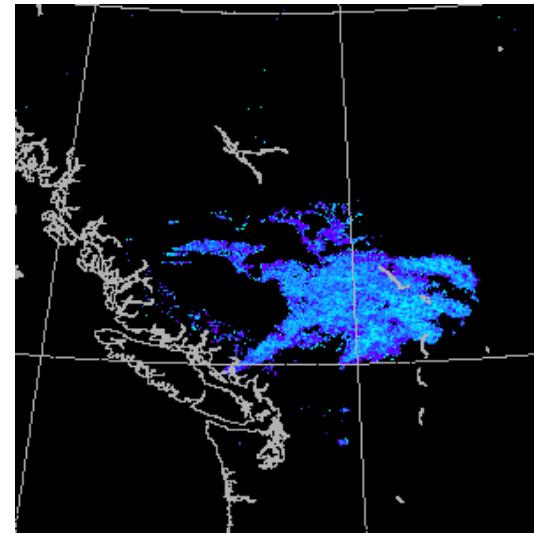


Retrieval of aerosol properties over western Canada on Aug. 14, 2018

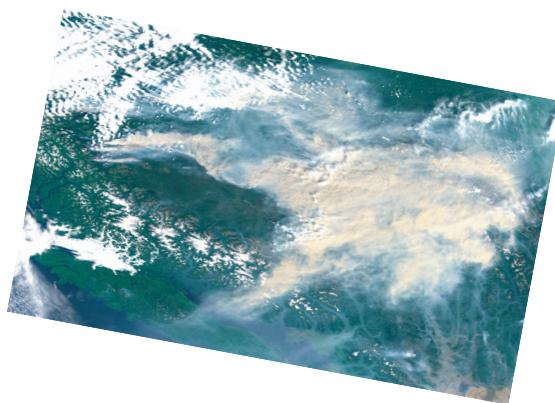
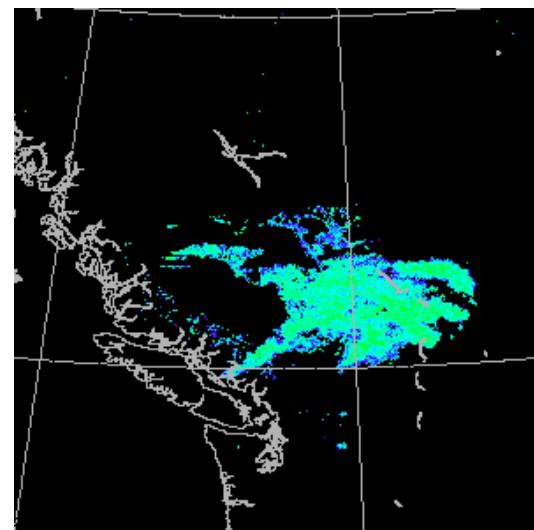
Ang Exp (AE)



SSA (380 nm)



SSA (530 nm)



b) Aerosol retrieval for finite atmosphere model

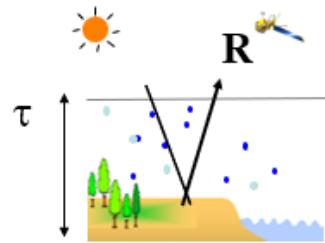
Inputs :

Retrieved CMF, complex refractive index (m) from VMSOS
Nadir view SGLI measurements at 380, and 530 nm.

Retrieved properties:

$AOT\ 550\ nm$

b) with boundary conditions for finite atmosphere

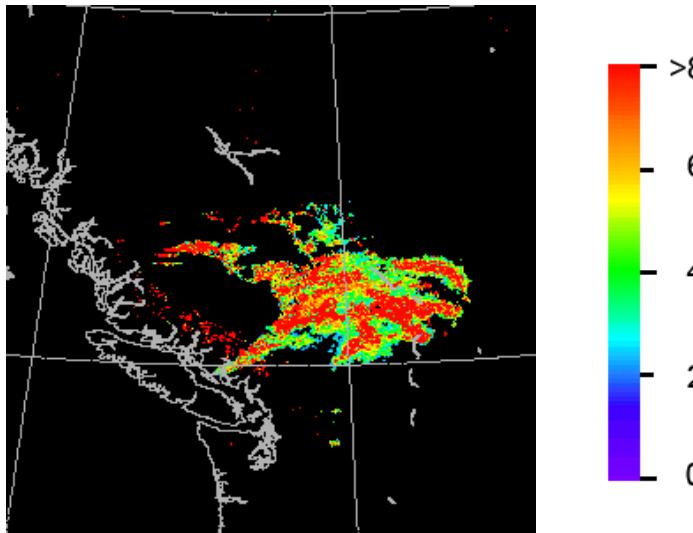


Outputs:

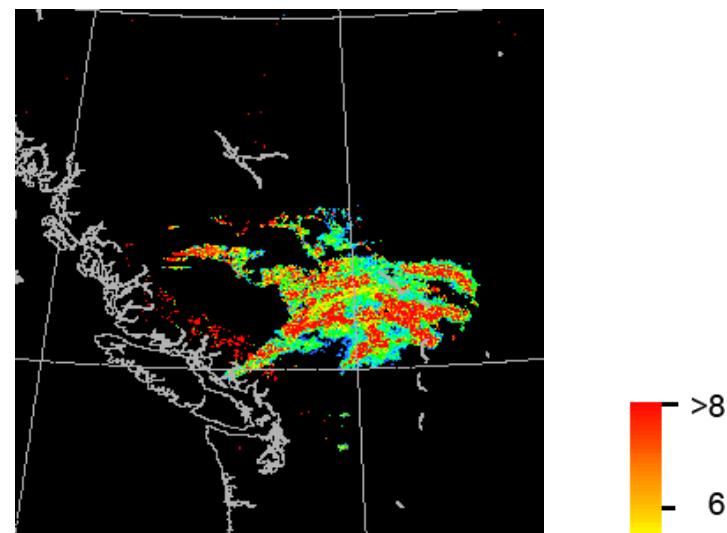
→ AOT (380, 412, 443, 530 nm)

Retrieval of aerosol properties over western Canada on Aug. 14, 2018

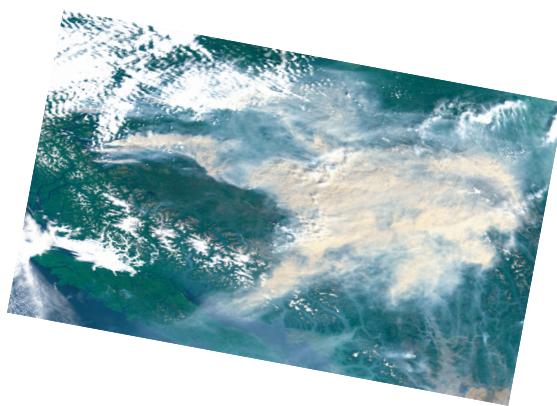
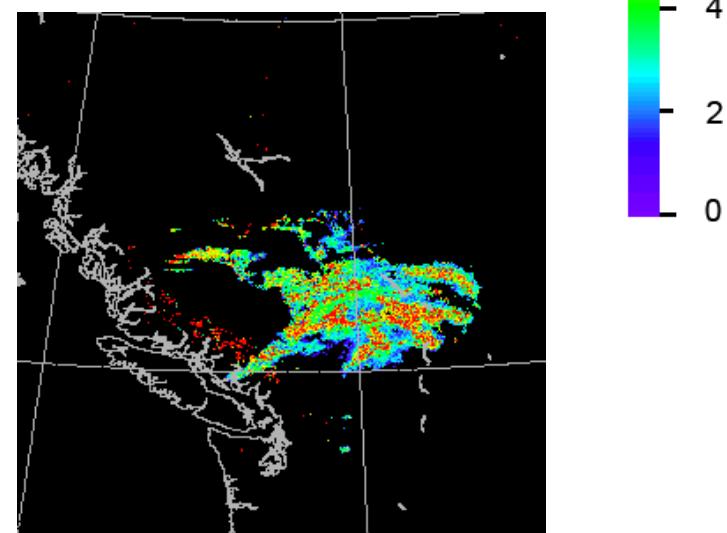
AOT (380 nm)



AOT (443 nm)



AOT (530 nm)



Summary

Detection and classification of absorbing aerosols based on color ratio of near UV (AAI, and DDI) is available for SG LI.

Radiation simulation algorithms are used for semi-infinite atmosphere model in the polarization field.

Aerosol properties as *Ångström exponent, SSA and AOT* are retrieved over dense aerosol region.

Further polarization information should be considered for efficient utilization.

Finally, I expect aerosol information from ground will be available by AERONET DRAGON campaigns.

Acknowledgements

The authors thank to JAXA for data distribution of GCOM-C/SG LI measurements.



BURGDORF