

# Droplet vertical sizing in shallow marine clouds using passive satellite measurements

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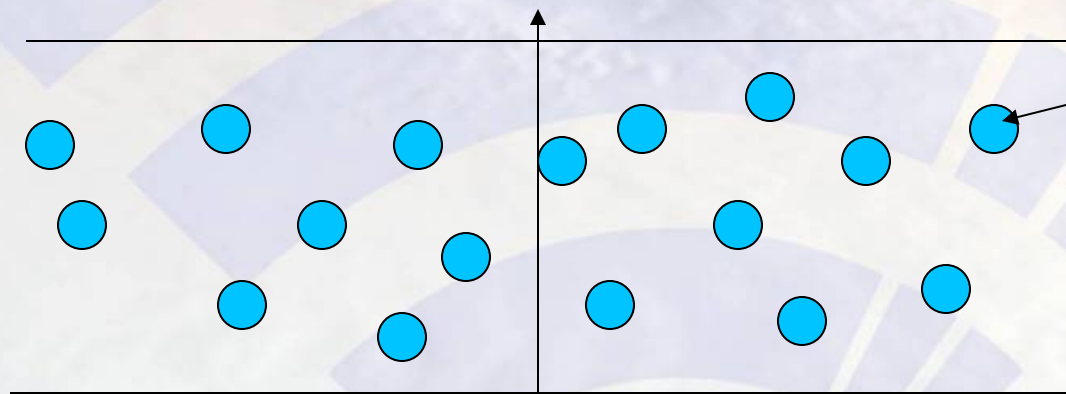
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# Contents

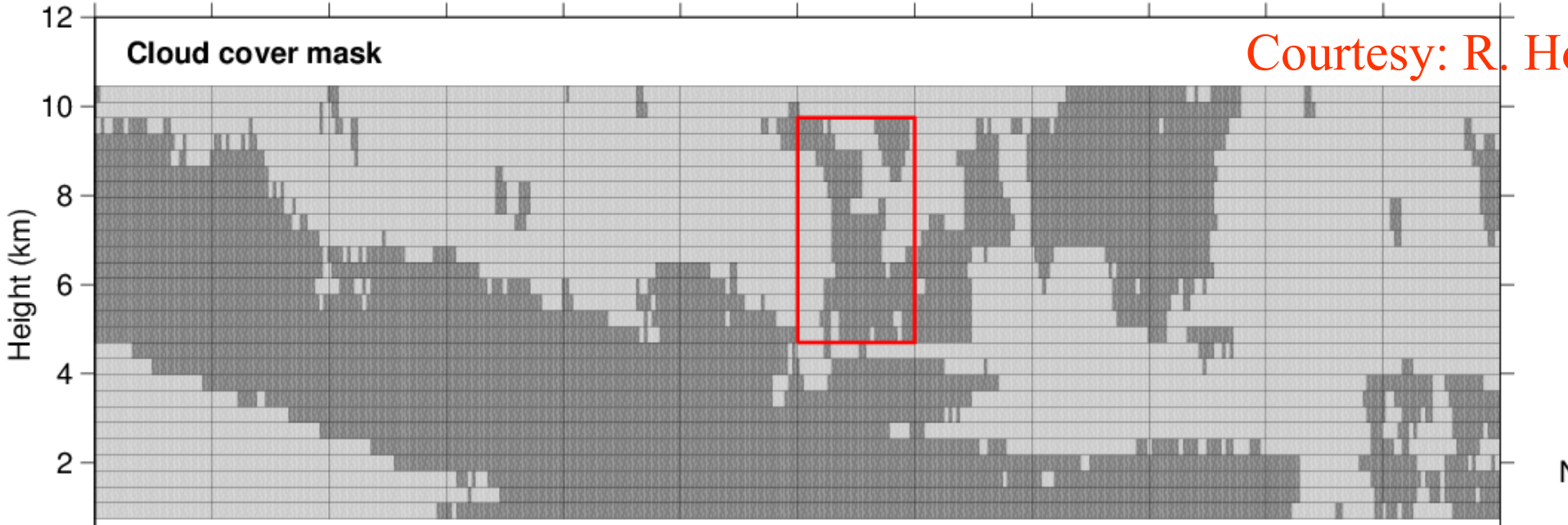
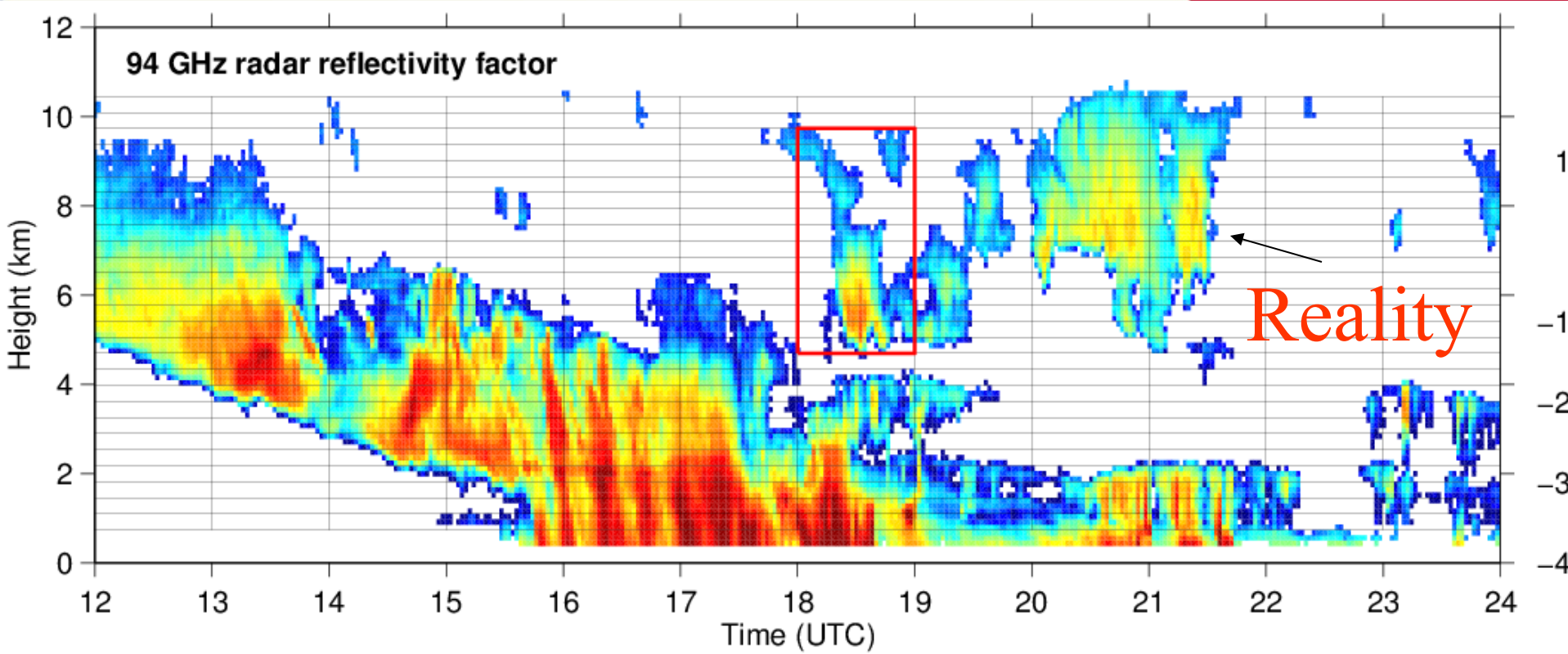
- Introduction
- Variation of spectral cloud reflectance
- Algorithm
- Application to synthetic, airborne and satellite data
- Conclusions

# 1. Introduction

Clouds are vertically inhomogeneous media but homogeneous cloud model is used in satellite cloud retrieval algorithms – contradiction.



Average size is the same on all levels  
In the clouds



Question: Is it possible to make profiling of water clouds using passive (and not exclusively active) observations?

Answer: **Yes – at least in some cases!**

**Idea:** Use of multi-spectral and multi-angular polarimetric hyperspectral observations (penetration depth difference for different probing wavelengths, observation geometries and polarization -----> sampling of different cloud volumes)

**References for use of MODIS observations:**

Chang and Li, 2002, 2003;  
Kokhanovsky and Rozanov, 2011

# 2. Theory

# Variation of reflectance

$$\delta R_{a_{ef}}(\lambda) = \int_0^1 \frac{\delta R(\lambda)}{\delta a_{ef}(z)} \left[ \lambda, \bar{a}_{ef}, z \right] \times \delta a_{ef}(z) dz$$

functional derivative

$z=Z/H$

$$\delta R_{a_{ef}}(\lambda) = \sum_{k=1}^{N_k} S_{a_{ef}}(\lambda, z_k) \frac{a_{ef}(z_k) - \bar{a}_{ef}(z_k)}{\bar{a}_{ef}(z)}$$

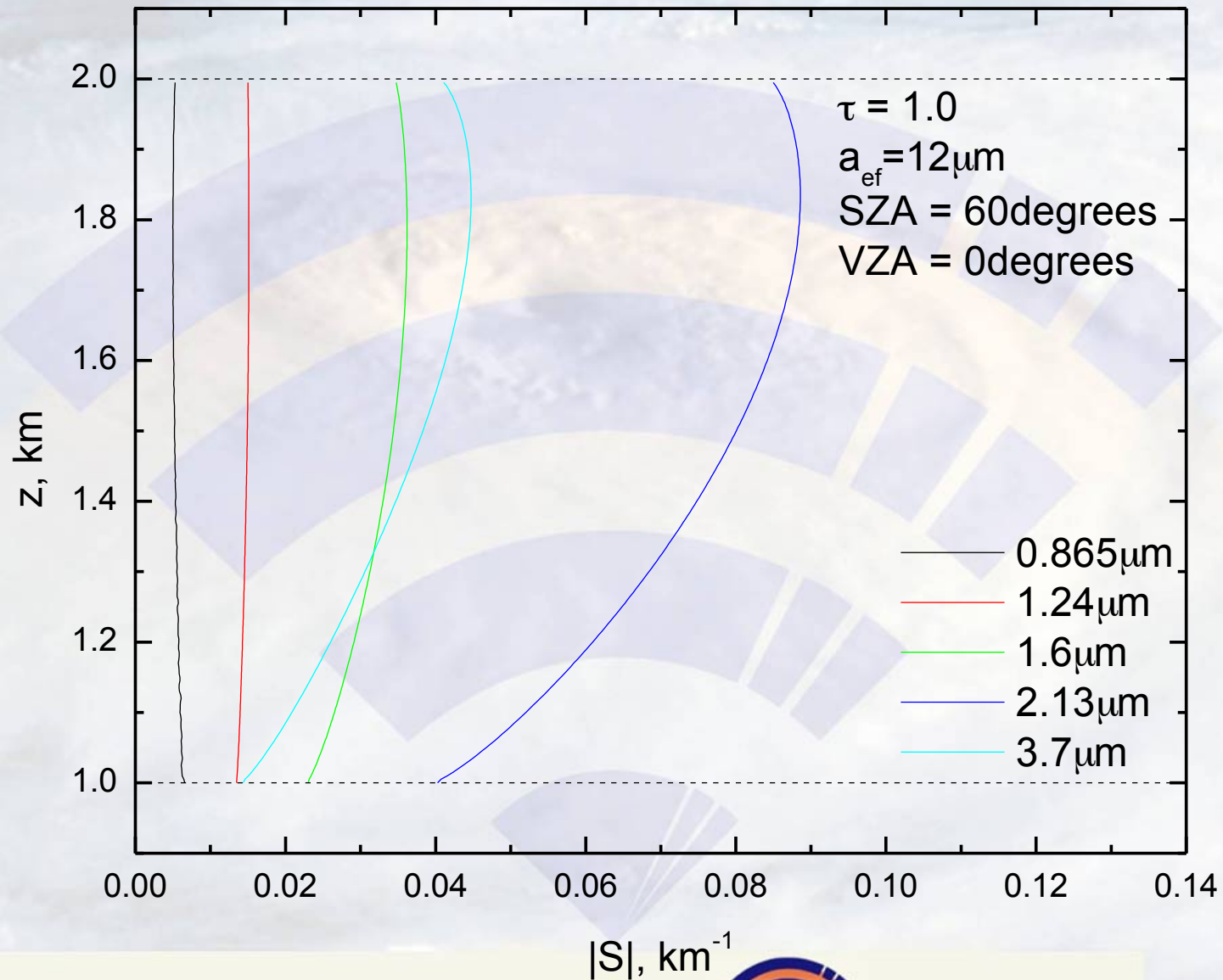
$$S_{a_{ef}}(\lambda, z_k) = \left( \frac{\delta R(\lambda)}{\delta \ln a_{ef}} \right)_k f_k \Delta z_k$$

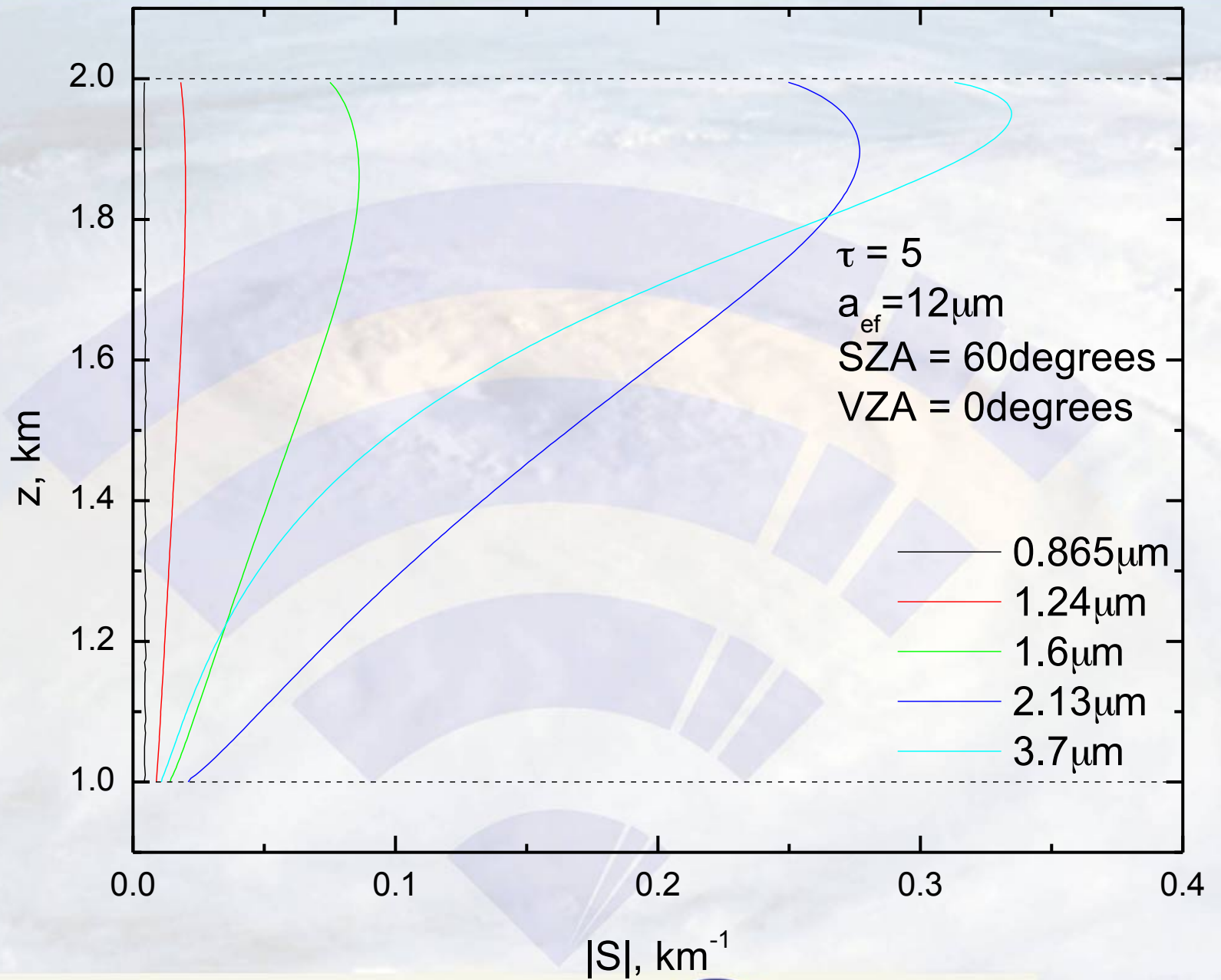
plots

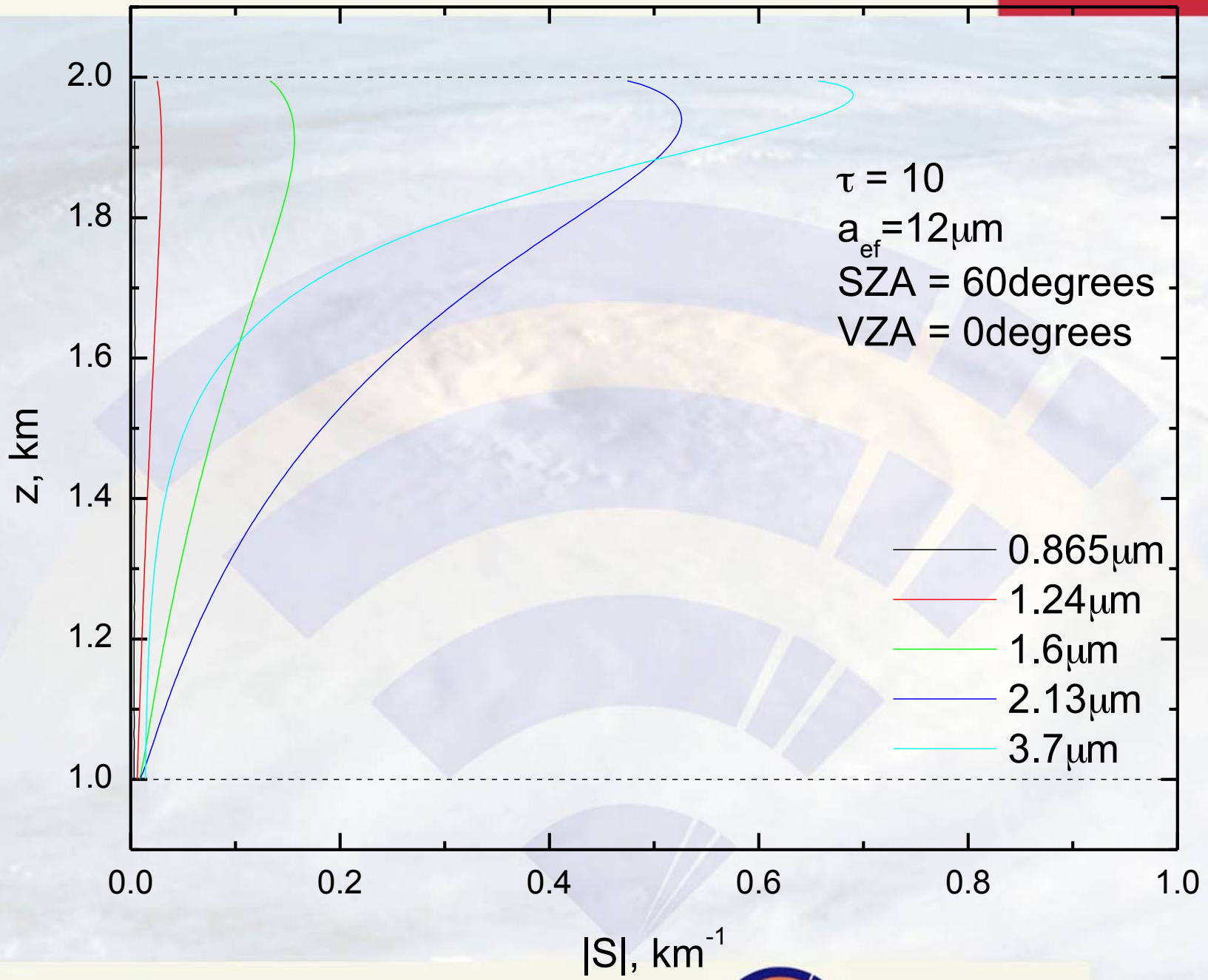
$\delta a_{ef}(z)$  variation of droplet size at depth  $z$

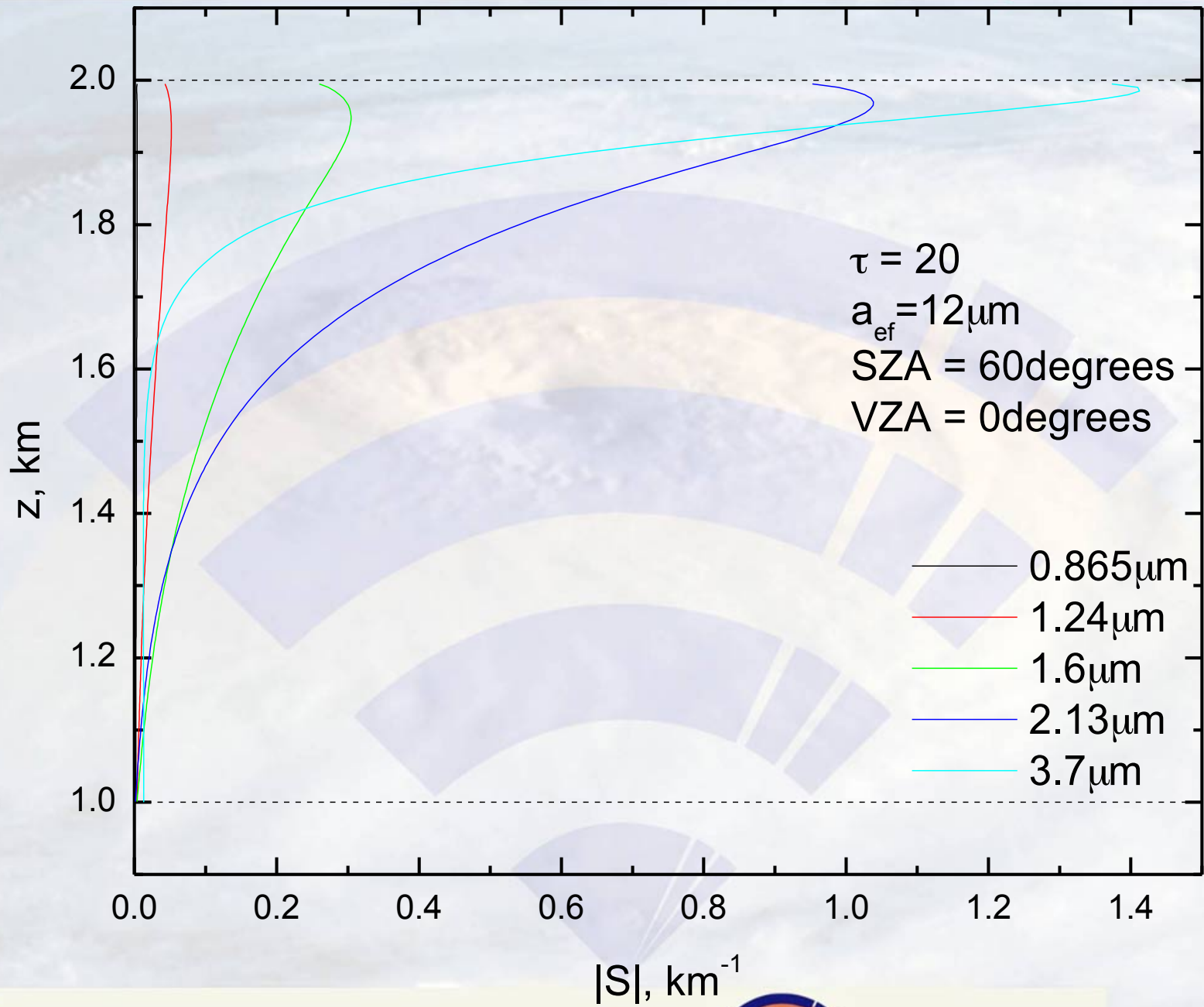


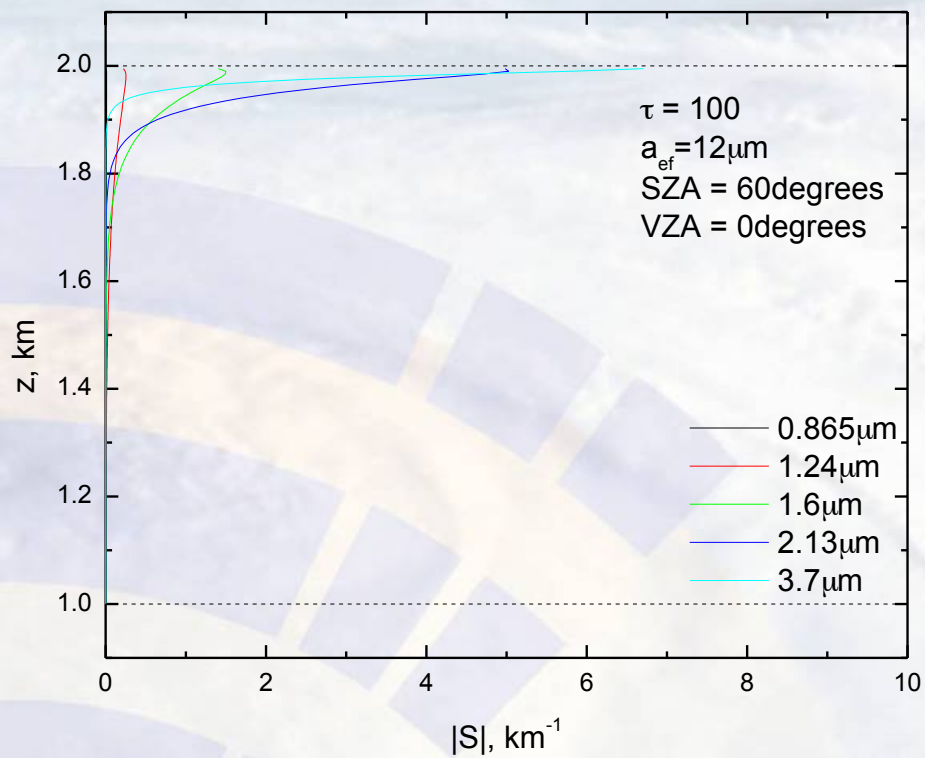
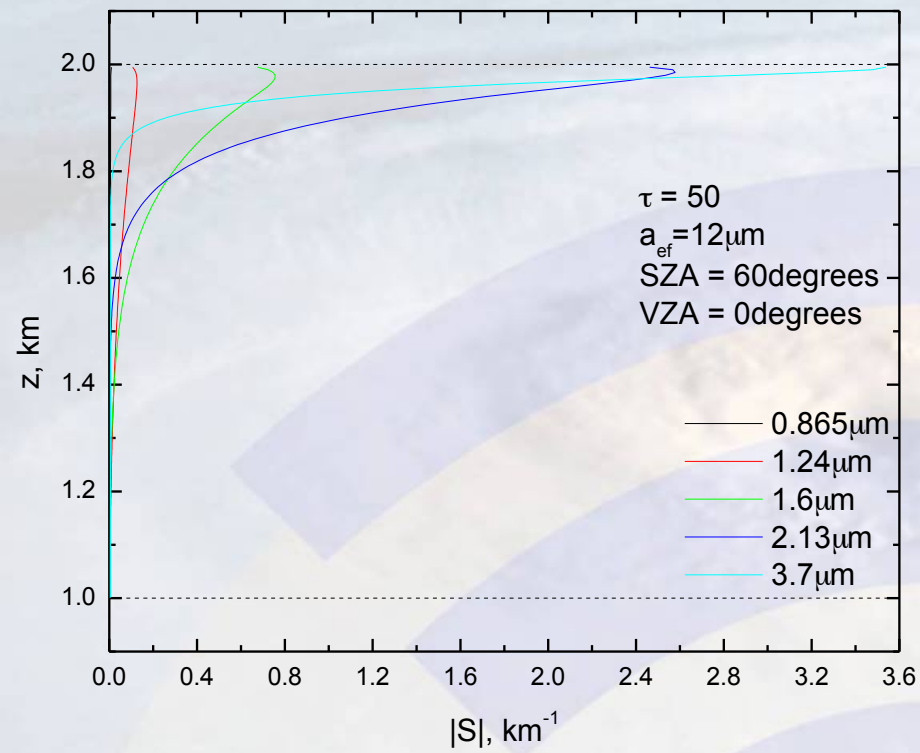
# Results of calculations











# Algorithm

$$\delta R(\lambda, a_{ef}(z), \overset{\mathbf{r}}{\Omega}) = \int_{z_1}^{z_2} W(z, \bar{a}_{ef}(z), \lambda, \overset{\mathbf{r}}{\Omega}) \delta a_{ef}(z) dz$$

$$a_{ef}(z) = A + zB$$

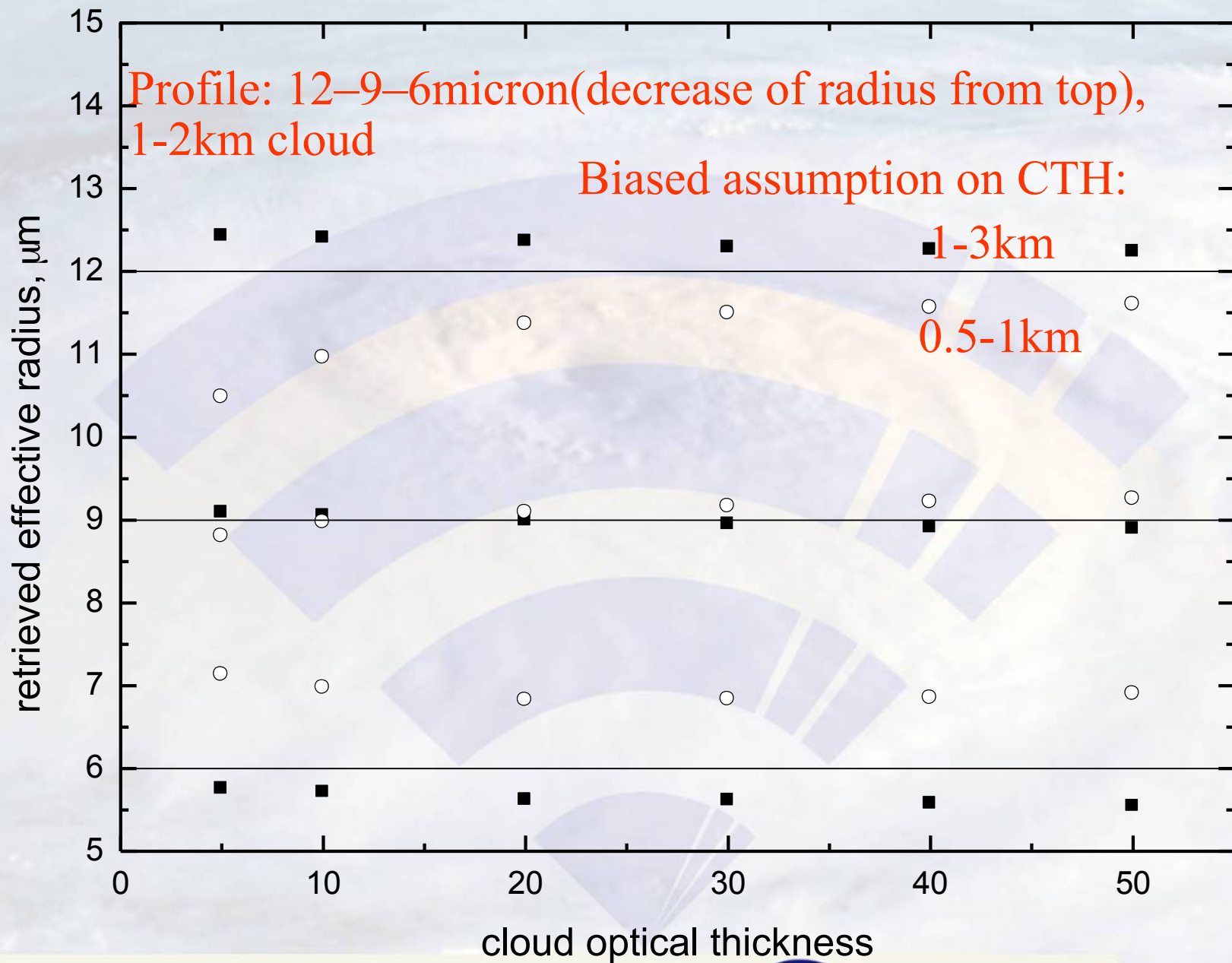
$$\delta a_{ef} = \delta A + z\delta B$$

$$\delta R(\lambda) = c_1(\lambda) \delta A + c_2(\lambda) \delta B$$

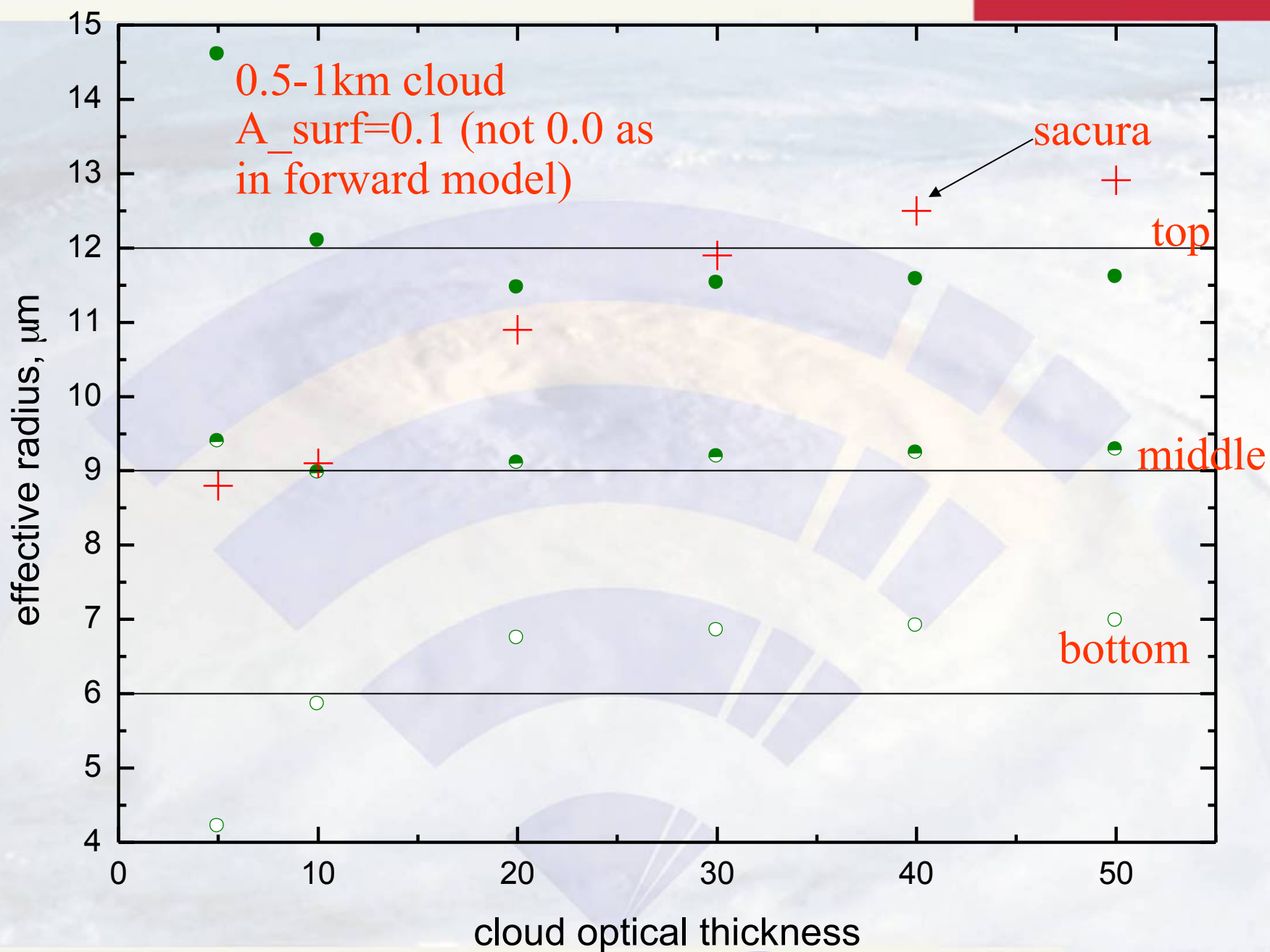
$$\overset{\mathbf{r}}{y} = \hat{c} \overset{\mathbf{r}}{x}, \quad \hat{c}^T \overset{\mathbf{r}}{y} = \hat{c}^T \hat{c} \overset{\mathbf{r}}{x}, \quad \overset{\mathbf{r}}{x} = [\hat{c}^T \hat{c}]^{-1} \hat{c}^T \overset{\mathbf{r}}{y}$$

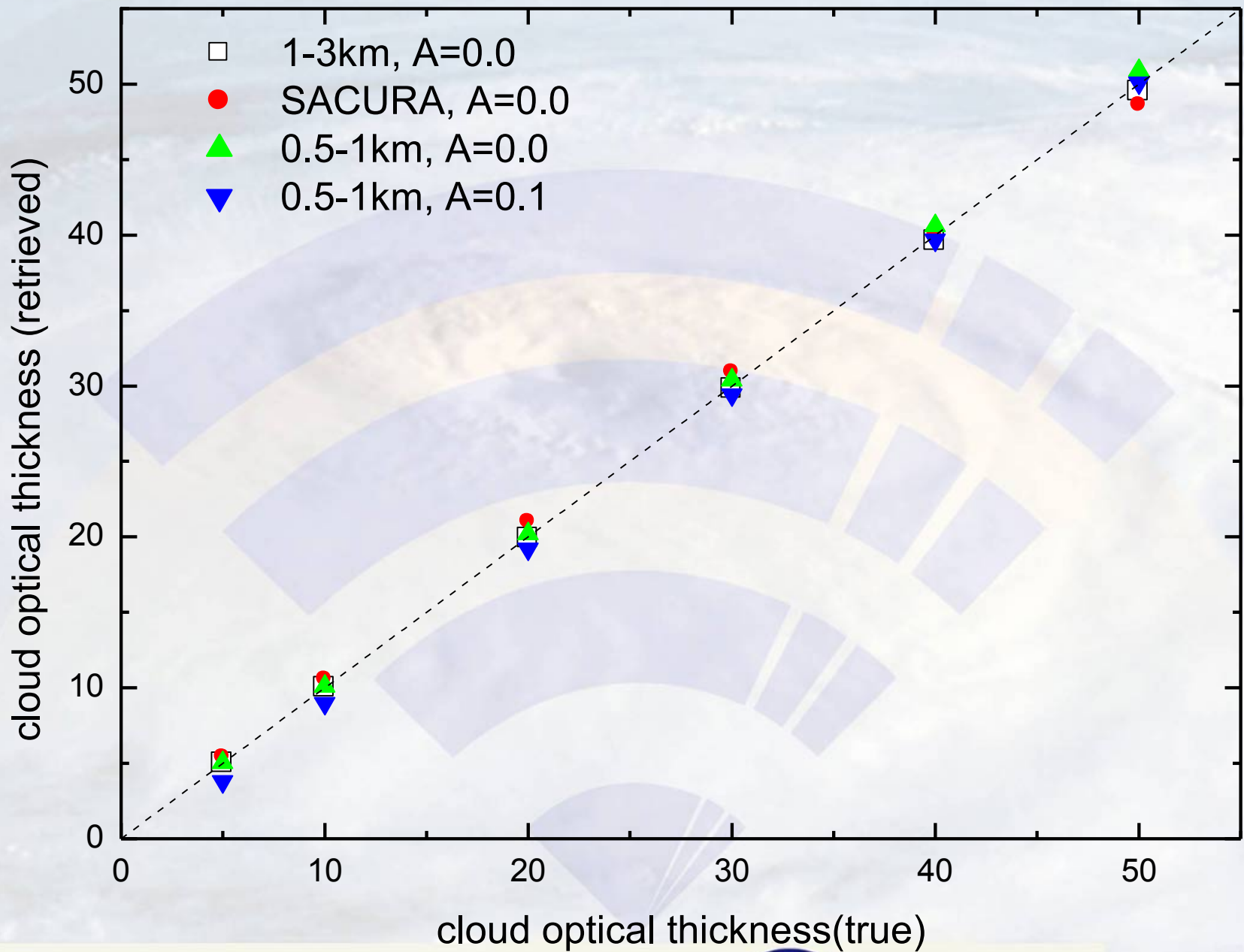
# 3. Retrievals

## 3.1 Synthetic retrievals

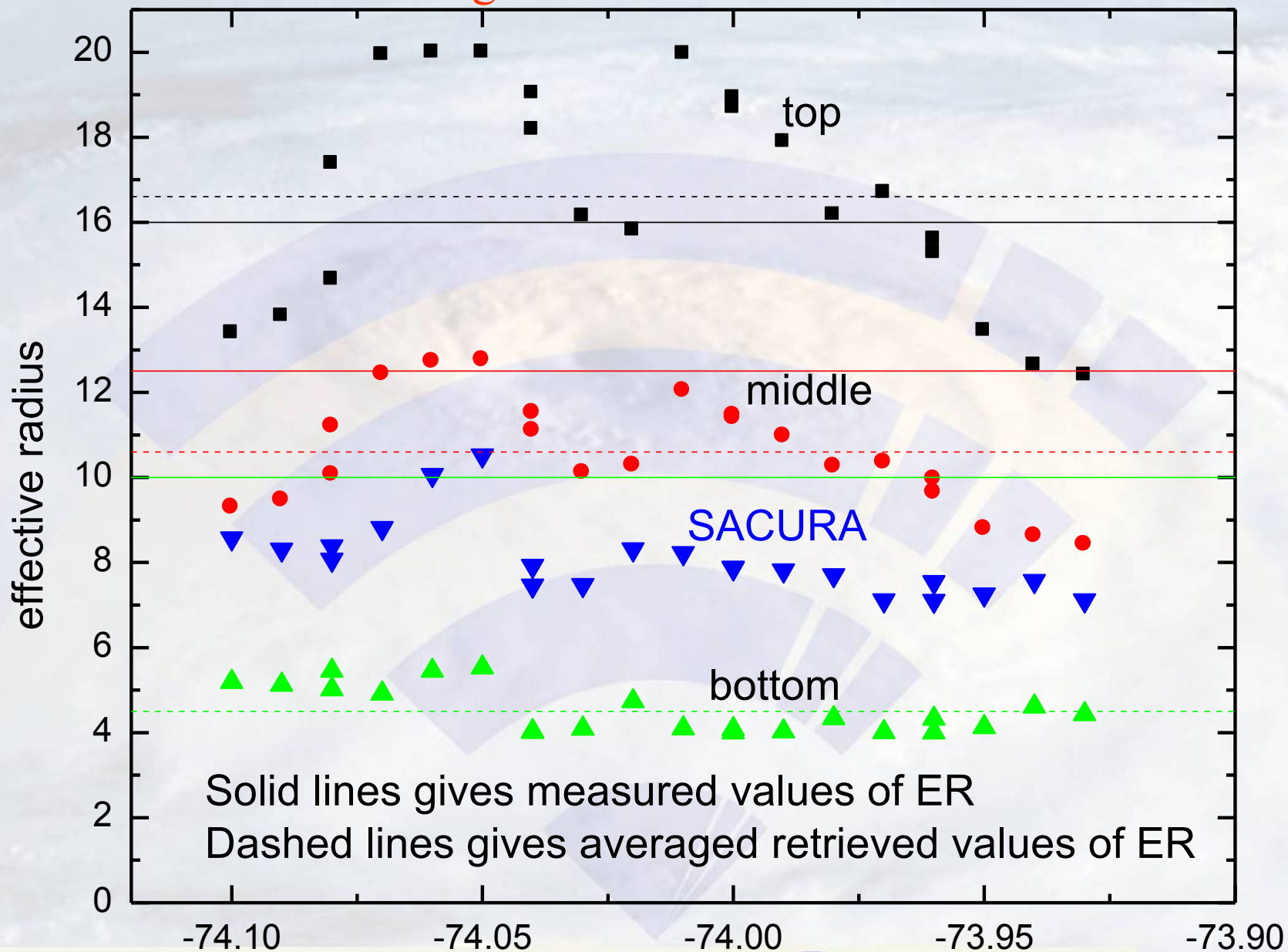








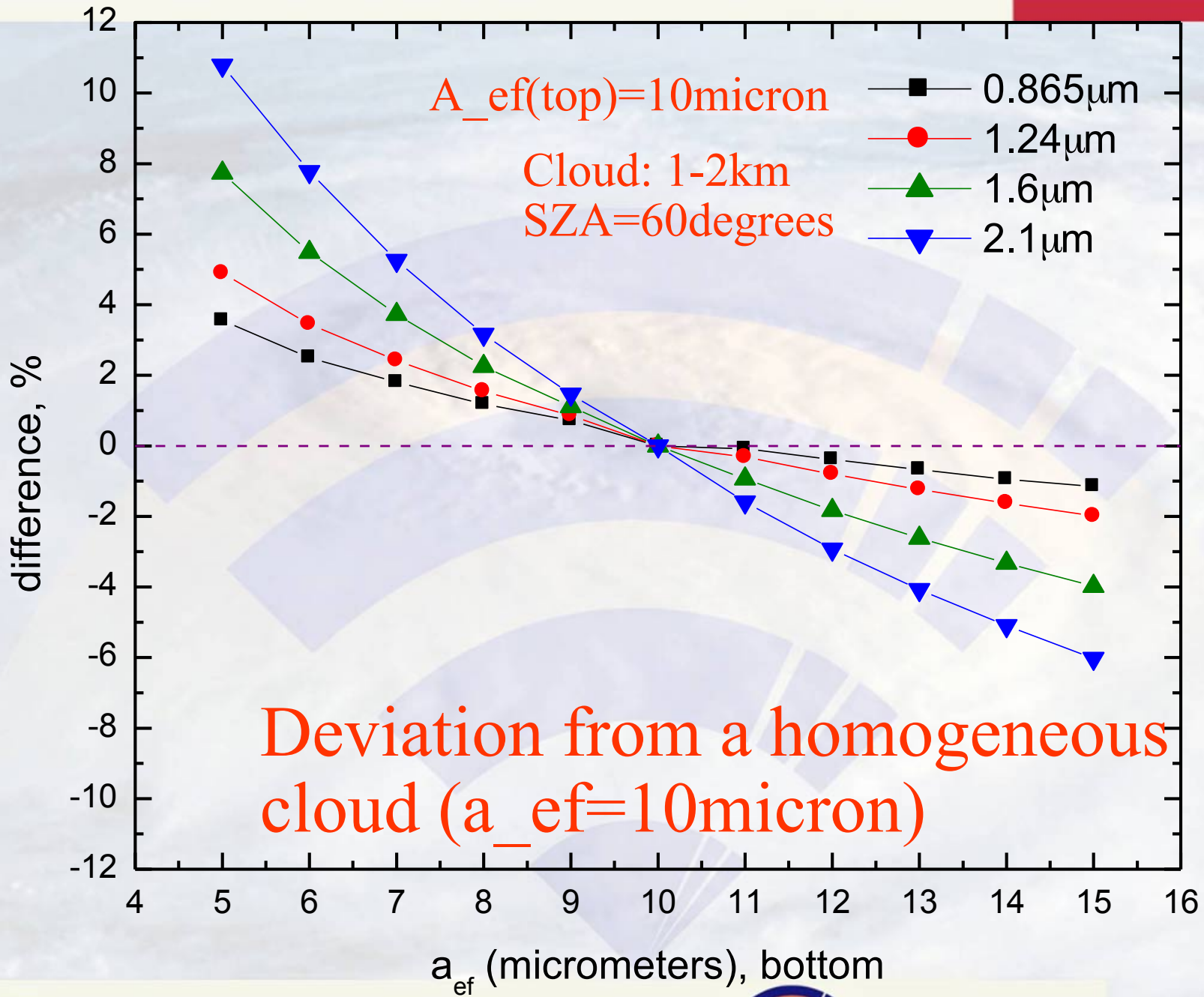
# Retrievals using aircraft data



Solid lines gives measured values of ER

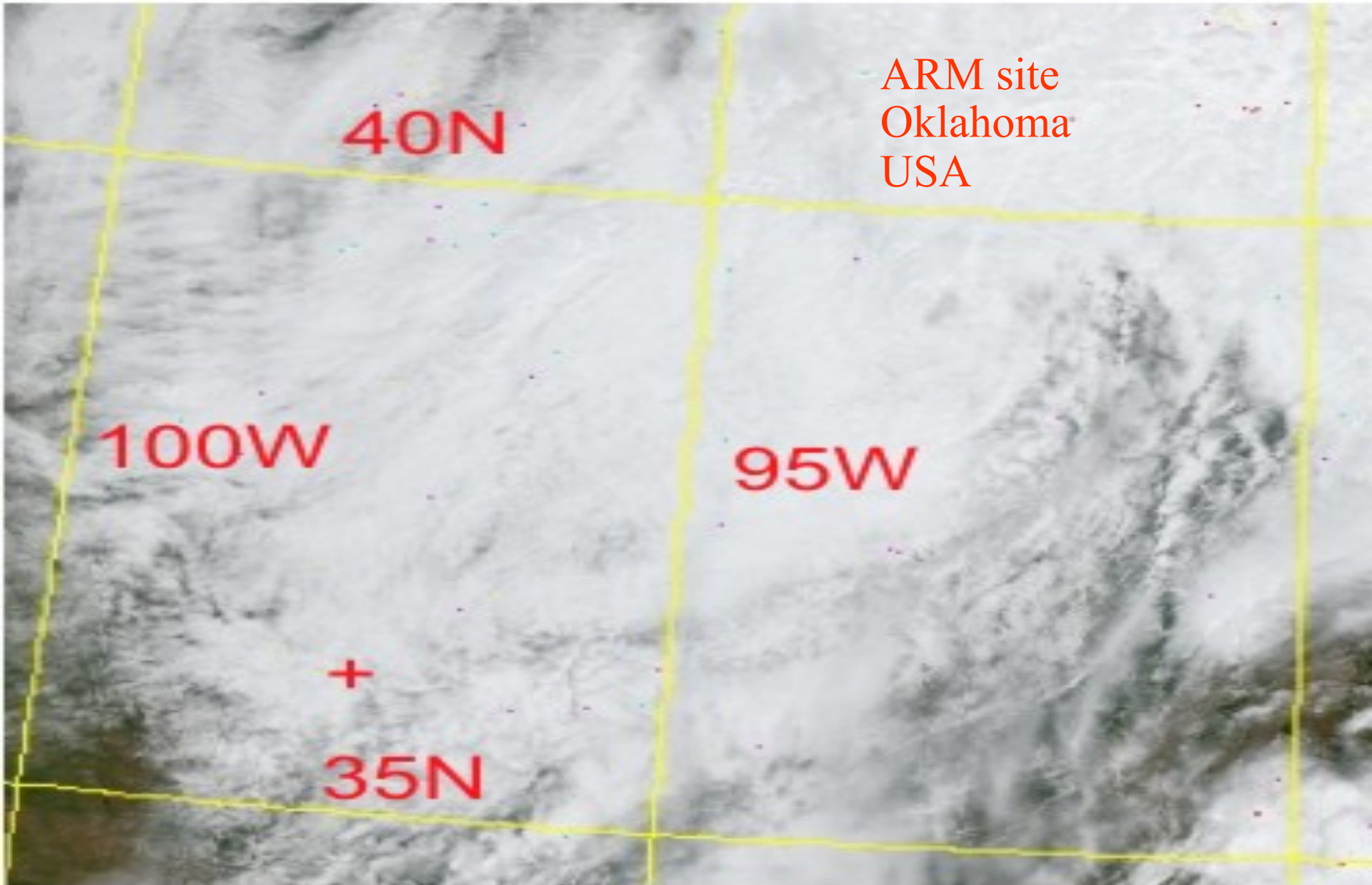
Dashed lines gives averaged retrieved values of ER

latitude, degrees



# Retrievals using MODIS data

## 1. Retrievals over land



40N

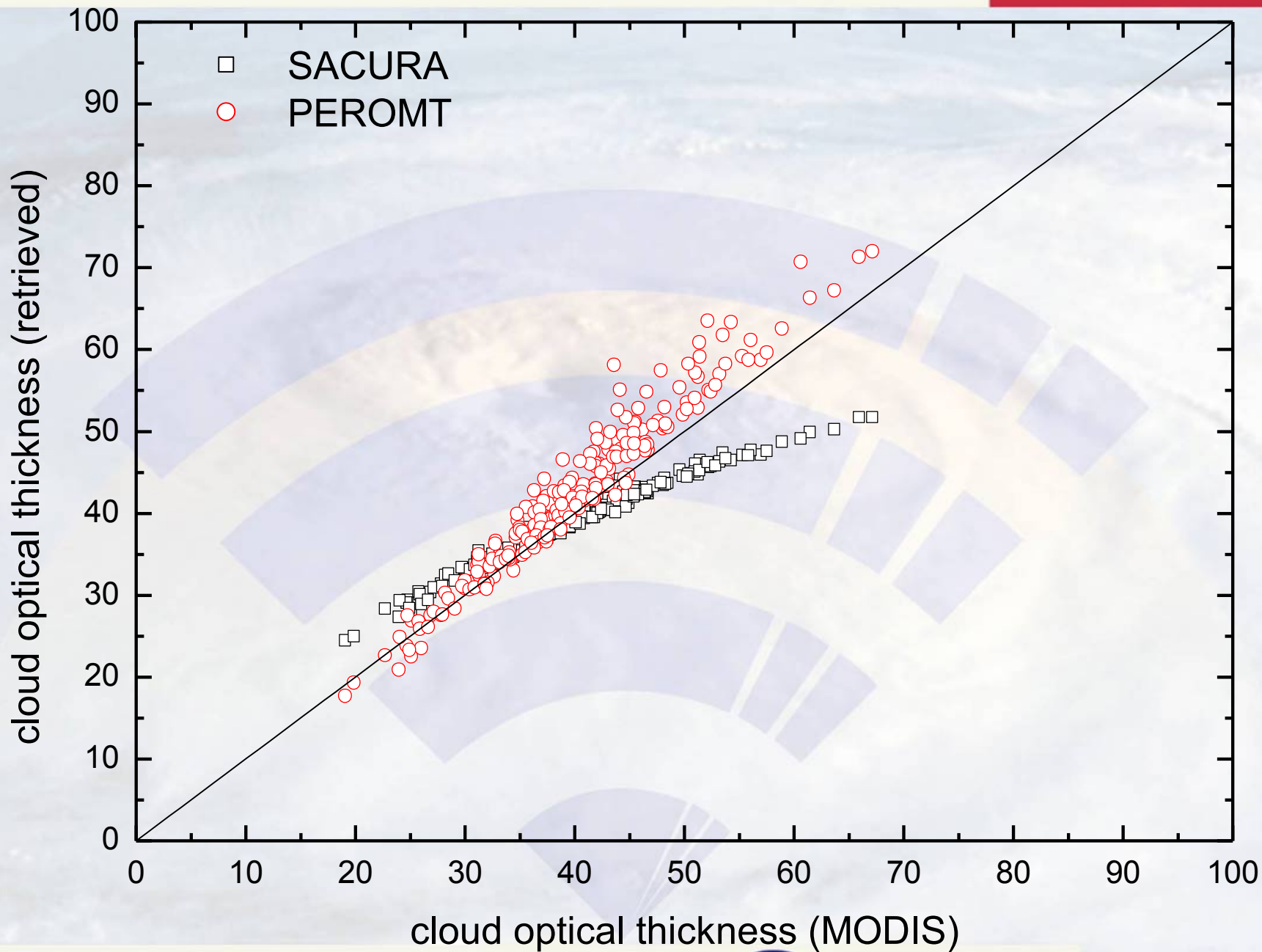
ARM site  
Oklahoma  
USA

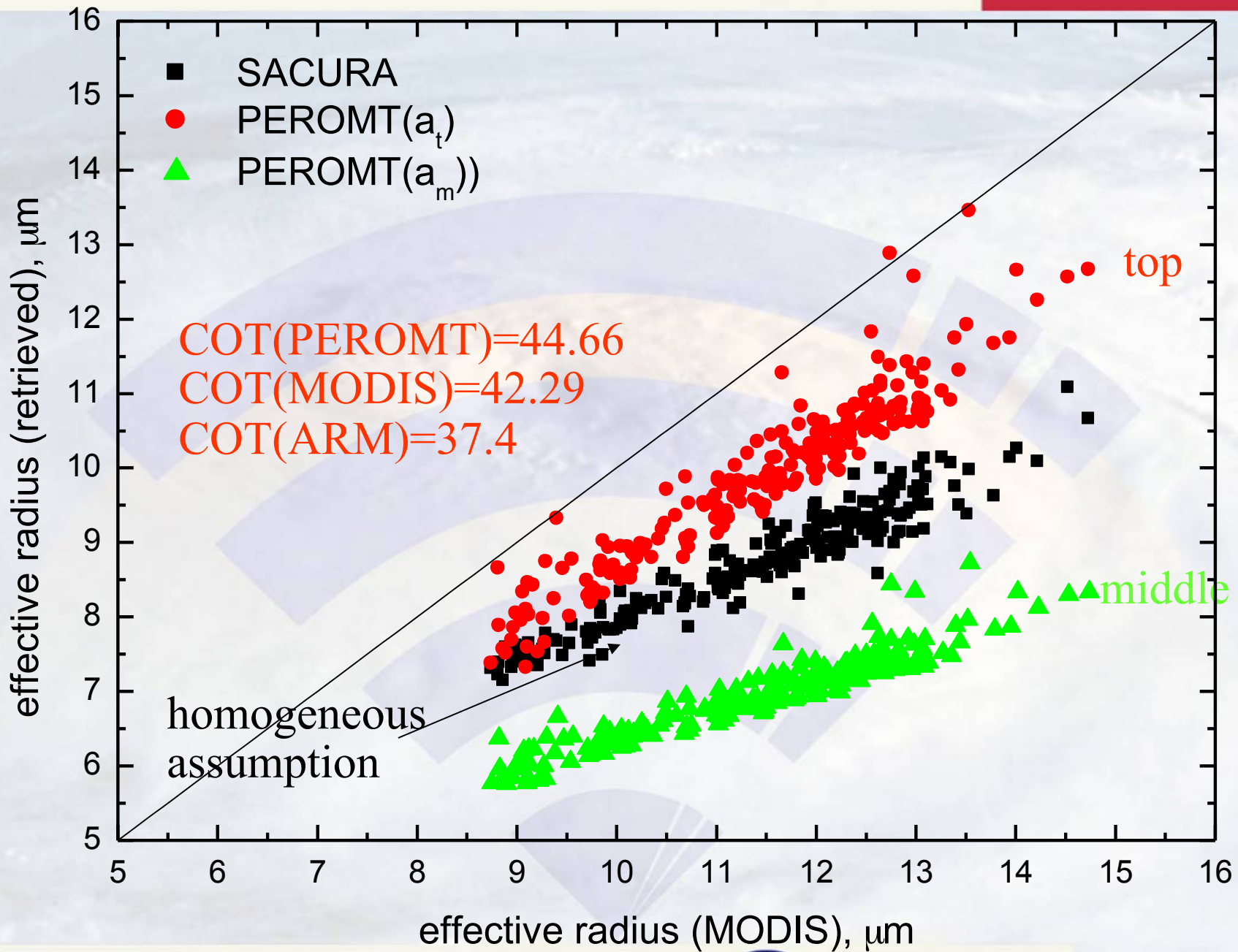
100W

95W

+

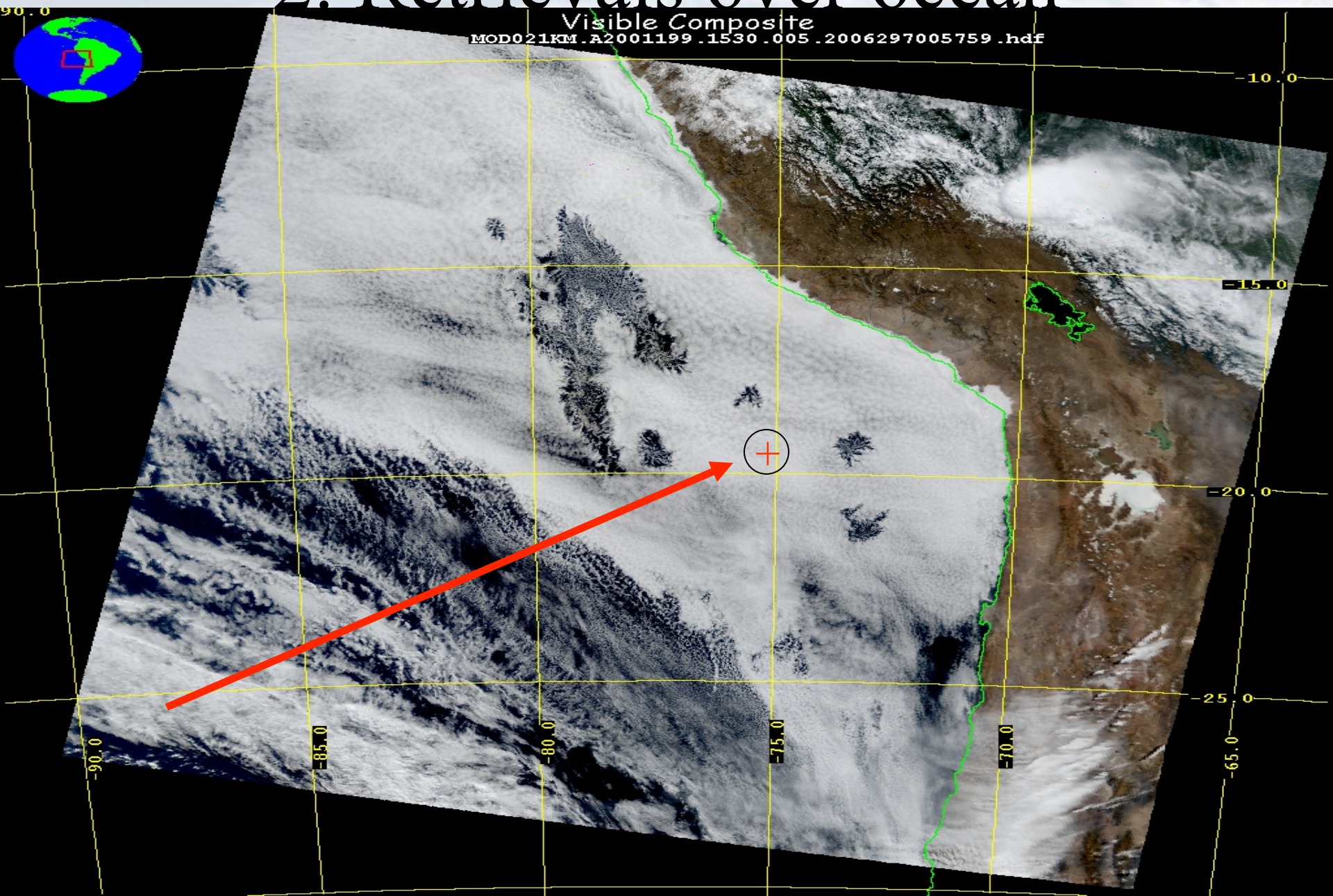
35N



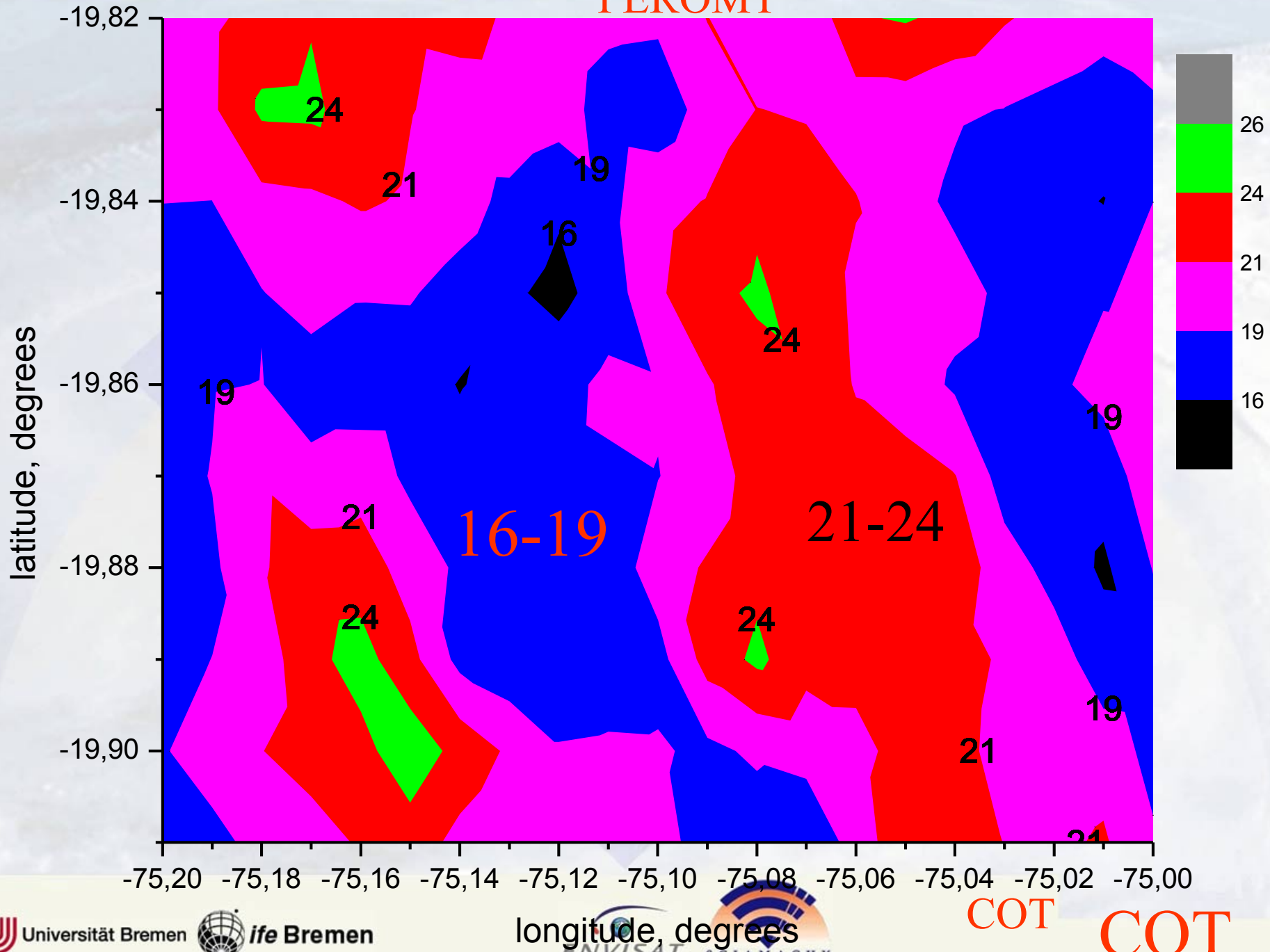




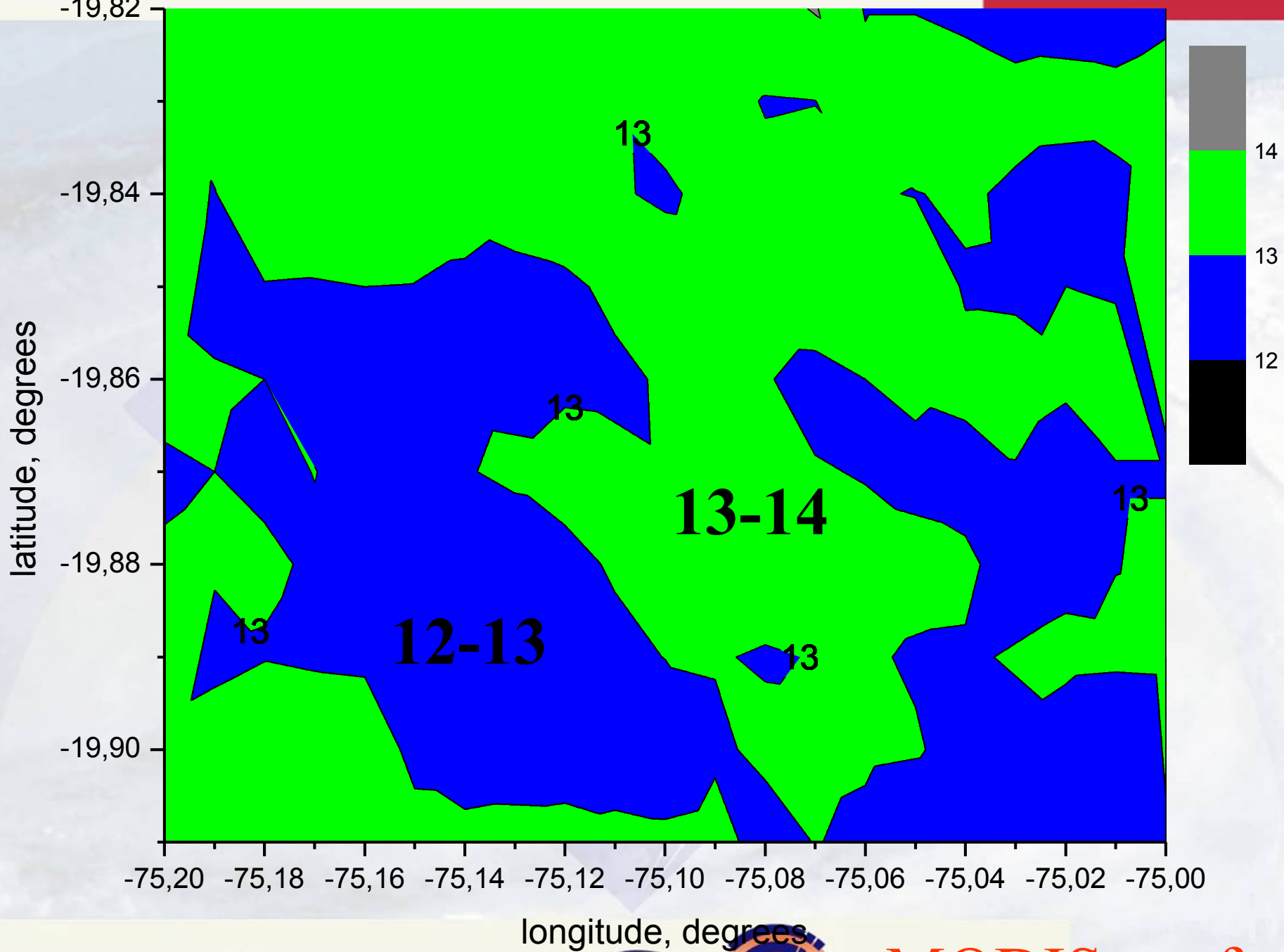
# 2. Retrievals over ocean



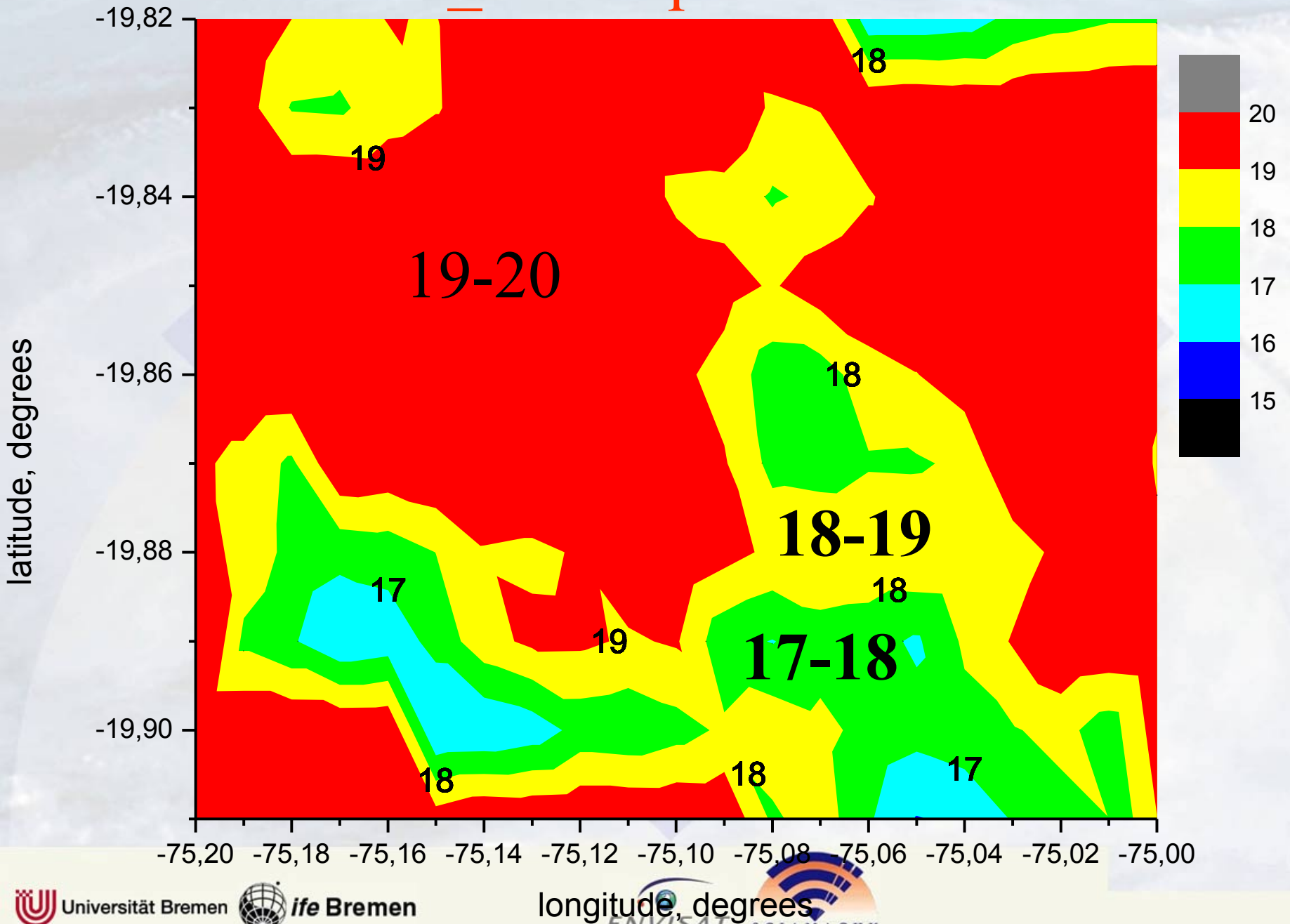
# PEROMT



COT COT

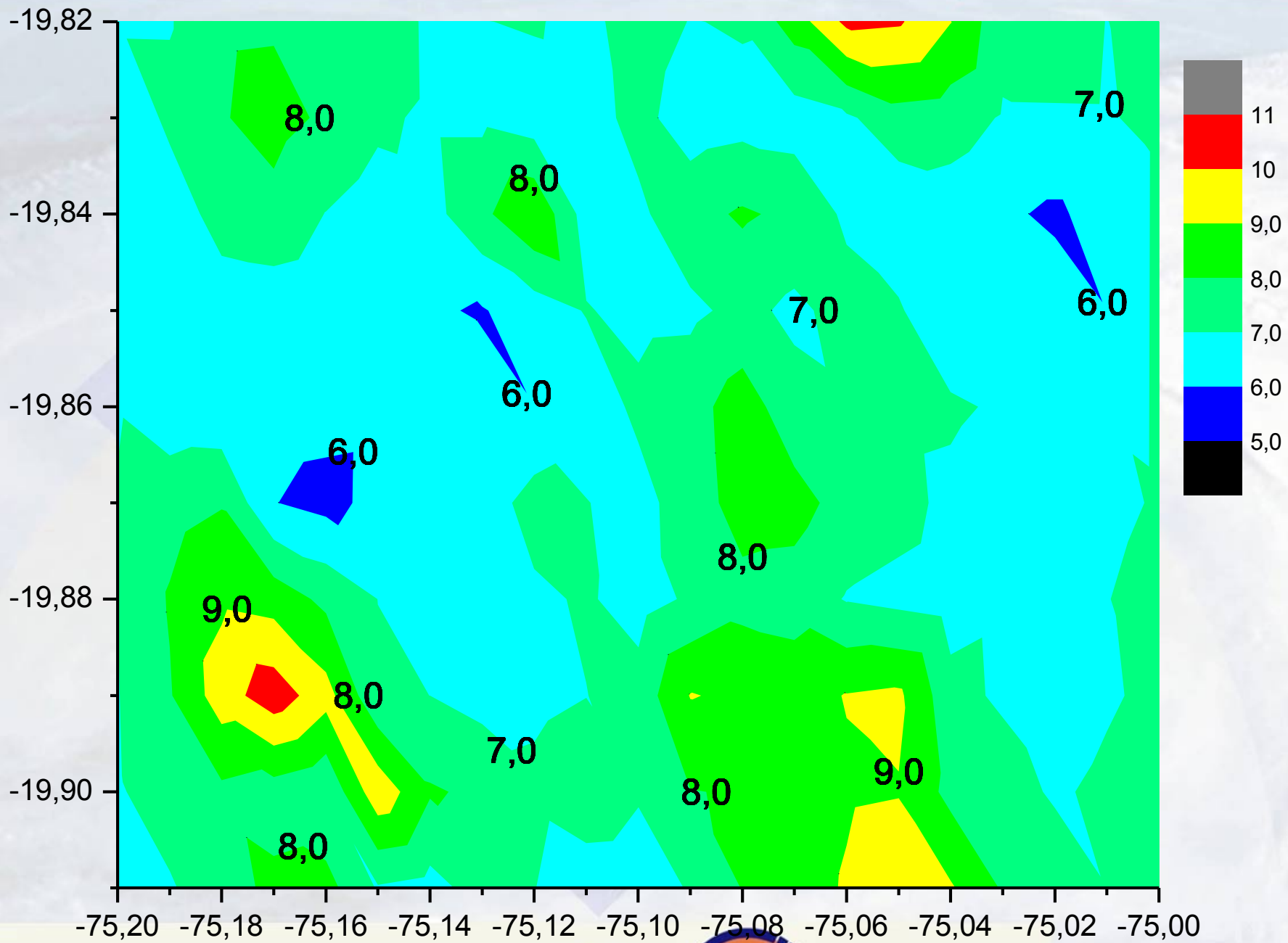


# a\_ef at top of cloud





latitude, degrees



-75,20 -75,18 -75,16 -75,14 -75,12 -75,10 -75,08 -75,06 -75,04 -75,02 -75,00

longitude, degrees

Bottom

# Conclusions and outlook

- It is possible to retrieve vertical profiles of effective radius of droplets in clouds using passive spectral measurements in some cases as demonstrated in the presentation. Further validation using in situ measurements is needed.
- The developed technique assumes that the cloud optical thickness is known from retrievals in the visible. Therefore, the problem is reduced to the case of finding a  $a_{\text{eff}}(z)$  because for a fixed assumed  $N(z)$  profile (up to a constant multiplier) and known COT, profiles  $a_{\text{eff}}(z)$  and  $N(z)$  are inter-related.
- The technique also enables the improvement of the liquid water path estimation

# Acknowledgements

- NASA
- ESA
- DFG
- Aircraft team
- VOCALS-UK team