aerosol climatology and applications



S. Kinne and colleagues

MPI-Meteorology

with support by AeroCom and AERONET

aerosol - a complex atmos. ingredient

- **aerosol** ('small atmos.particles')
 - many sources
 - short lifetime
 - diff. magnitudes in size
 - changing over time
 - aerosol $\leftarrow \rightarrow$ clouds
 - aerosol $\leftarrow \rightarrow$ chemistry
 - aerosol $\leftarrow \rightarrow$ biosphere
 - aerosol ←→ aerosol

highly variable in space and time !





simplify !

- aerosol modules can be (very) expensive
 - stratification by type and size
 - mixing states, phase transitions, chemistry
 - environmental processing (e.g. clouds)
 ... still important processes may be missed
- 'climate applications' demand simplifications
 what are the optical properties ?

 ... to quantify (direct) radiative impacts
 what are is CCN / IN concentrations ?

 ... as a link to (indirect) radiative impacts

past simplifications

optical properties

- Tanre CLIMATOLOGY (1984 !)
 - dust 'bubble' over Africa RF(1.55, 0.005)
 - \rightarrow 20 W/m2 too much solar absorption
- Koepke GADS / OPAC (1997)
 - -10 aerosol types \rightarrow mixtures allowed
 - -size (&opt. prop) function of relative humidity
 - -only for jan and july, 5x5 resolution

CCN estimates

Nakajima AI ~ CCN (2001)
 AI = AOD * Angstrom

revisit climatologies

better and more observations

- 10 + years sun-/sky-photometer networks
 statistics from ca 500 sites
- 10 + years dedicated satellite sensors
 MODIS, MISR, POLDER, OMI, ATSR, CALIPSO

advancements in modeling

- detailed aerosol component in global modeling
 AeroCom, ...
- kappa method (Petters, Kreidenweiss)
 - supersat. & kappa \rightarrow crit. radius \rightarrow CCN / IN

start with modeling

ensemble-median

- advantages
 - complete
 - consistent
 - 'tuned' to satellite patterns

issues

as good as
assumptions

all relevant aerosol optical properties

- aerosol optical depth (AOD)
- single scattering albedo
- Angstrom (→ asymmetry factor)

annual aerosol optical depth at 550nm



improve with AERONET data

AERONET

advantages high quality

- issues
 - local
 - uneven distr.

NO satellite data or in-situ data are directly involved



model improvements

- merging for 3 mid-visible aerosol properties
 - **AOD** aerosol opt. depth ... info on amount
 - SSA single scatter. albedo ... info on abs.
 - Angstrom parameter ... info on size • Ap
- improvements indicate general deficiencies in global aerosol component modeling
 - lack of AOD (e.g. SH biomass burning)
 - missing absorption
 - overestimate size

- (e.g. biomass / pollution)
 - (e.g. boreal biomass)

... next: comparison of seasonal maps

AOD_550nm

model / merged / AERONET



SSA at 550nm

model / merged / AERONET



Ap at 550nm

model / merged / AERONET



spectral / temp. extension

- use Angstrom parameter to split AOD in
 - AOD, f (fine-mode particles > 0.5μm radius)
 - AOD, c (coarse-mode ... part. < 0.5μm radius)
- use modeling to split fine mode AOD (AOD, f) in
 - AOD, f_pre-industrial (AOD,f yr.1850)
 - AOD, f_anthropogenic (AOD,f AOD,f yr.1850)
- temporal changes ... only for AOD, f_anthrop.
 - coarse mode ... no change with time (BAD !)
 - pre-industrial fine ... no change with time
 - anthropogenic fine ... changes ($1860 \rightarrow 2100$)



optimistic forcing futures

- translate temporal changing anthropogenic AOD maps into direct radiative forcing
 - global, annual, at ToA, for all-sky conditions
 - year 1860 0.0 W/m2
 - year 1950
- 0.2 W/m2
- year 2000
- 0.5 W/m2
- year 2100 - 0.3 W/m2 (all three RCPs !!) -very optimistic views on future direct effects
- anthropogenic is small compared to natural
 - anthropogenic (today)
 - natural (corase & pre-ind) 1.2W/m2
 - volcanic stratosph. (1992)
- 0.5W/m2
- 2.0 W/m2

future anthrop – ToA forcing to reduce by 2100 to about 65% of today's



what about the indirect effect ?

 an aerosol climatology is more useful with direct connections to cloud- interactions

– can we associate aerosol anthropogenic change with anthropogenic changes in CCN ?

– can we provide estimates for IN ?

associated CCN / IN (1)

- AOD → concentrations
 - use altitude distr. to determine extinction (/m)
 - use different altitude distributions for the coarse and fine mode (by modeling / CALIPSO)
 - establish log-normal size-distr. parameters
 - fine-mode

-std dev.: 1.7 (defines the distribution width) $r = \frac{1}{2} \left(\frac{1}{2} - \frac{1}{2} \right)^2$

- $-\mathbf{r} \mod \mathbf{\leftarrow} \mathbf{Ap}$, fine (defined via 'fine' Angstrom)
- coarse mode is (pre-)defined (sea-salt or dust)
 sea-salt: 1.25/1.7 (r mode [μm] / std dev.)

-dust: 0.93/1.55 (1.23/1.7 1.56/1.85 1.95/2.00 2.31/2.15)

- assume horizontal homogeneity (1/m \rightarrow 1/m3)

associated CCN / IN (2)

assumptions

- coarse mode concentrations: all CCN
- coarse mode dust concentration: IN
- fine-mode concentrations: fraction is CCN
 - only concentrations larger than a critical radius are considered to be CCN
- anthropogenic (fine-mode) AOD fractions stratify fine-mode CCN into
 - anthropogenic (fine-mode) CCN ... and
 - pre-industrial fine-mode CCN ...
 - natural CCN (= pre-indust.fine +coarse)

associated CCN / IN (3) to reduce by 2100 to about 65% of today's

- fine-mode critical cut-off size
 - function of super-saturation [0.05 ... 0.20%]
 - function of temperature
 - function of 'kappa' (humidification factor)
 - over oceans near 0.7
 - over continents near 0.4
 - global monthly 1x1 maps of kappa are based on aerosol fine-mode component mass simulations for current aerosol conditions

- kappa = (1.0*mSS + 0.6*mSU + 0.1*mOC) /
 (mSS +mSU +mOC +mBC +mDU)

– compare current nat.CCN / anthrop CCN / IN

CCN at 1.5km above ground & IN (ice nuclei) anthropogenic CCN concentrations 'dominate'

natural CCN

anthrop.CCN

(dust) IN



summary

- a new 1x1 monthly aerosol climatology (yr.2000)
 - simplified and reasonable ... using new INFO
 - all needed opt.properties: AOD, SSA, g \dots (λ)
- extension in time (anthropogenic change)
 - very general for the direct impact, as the metinfluence on natural aerosol is not considered
 - exciting link for the indirect impact, as CCN are extracted and strongly linked to fine-mode
- potential improvements
 - better altitude definition by replacing model estimates with CALIPSO statistics
 - fine-mode dependence on relative humidity
 - better than 1x1 spatial resolution





... a longer version

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highly variable in space and time !





aerosol – its uneven spatial distribution

combining MODIS, MISR, AVHRR data



aerosol – its uneven vertical distribution



most of the tropospheric aerosol is near the surface

aerosol and climate ?

- aerosol (direct) impact on the energy balance
 - reflection of solar energy to space
 - absorption of solar energy (dust, BC, OC)
 - enhancement of IR greenhouse (elev. dust)
- aerosol impact is small compared to cloud imp.

why then all the interest in aerosol?

- some of the aerosol in anthropogenic added
 - anthrop. aerosol dominates in some regions !
 - aerosol can serve as nuclei for cloud and can affect clouds, precipitation and hydro.cycle !

simulating aerosol = \$\$\$

- complex aerosol modules exist (e.g. Sprintars)
- they are useful ... but often too expensive

the ECHAM/HAM aerosol modules carries 27 tracers !

optically active and relevant sizes

MODES IN M7	SOLUBLE/MIXED	INSOLUBLE
NUCLEATION (r < 0.005 µm)	1 N ₁ , M ¹ _{SO4}	
AITKEN (0.005 μm < r < 0.05 μm)	2 N ₂ , M ² _{SO4} , M ⁵ _{BC} , M ⁹ _{OC}	5 N ₅ , M ⁸ _{BC} , M ¹² _{OC}
ACCUMULATION (0.05 μm < r < 0.5 μm)	3 N ₃ , M ³ _{SO4} , M ⁶ _{BC} , M ¹⁰ _{OC} , M ¹³ _{SS} , M ¹⁵ _{DU}	6 N ₆ , M ¹⁷ _{DU}
COARSE (0.5 μm < r)	4 N ₄ , M^4_{SD4} , M^7_{BC} , M^{11}_{OC} , M^{14}_{SS} , M^{16}_{DU}	7 N ₇ , M ¹⁸ _{DU}

- Two-moment aerosol scheme that predicts mass and number concentrations of sulfate (SO4), black carbon (BC), particulate organic matter (OC), mineral dust (DU) and sea salt (SS)
- ► → Total of 27 prognostic equations

Hamburg Aerosol Module (HAM) [Stier et al., ACP, 2005]

how to simplify ?

... for tropospheric aerosol

- use an aerosol **climatology**, which gives
 - aerosol radiative properties
 - AOD, SSA, g ... at all needed wavelengths
 - separately for coarse and fine mode
 - aerosol AOD anthropogenic fraction (time)
 - fraction of fine mode than is anthropogenic
 - aerosol altitude distribution
 - CCN concentration
 - to maintain a link to clouds

climatology

- monthly 1x1 (lat/lon) resolution
 - compose one typical current year (year 2000)
 - median of 15 AeroCom models (no extremes)
 - enhance with quality AERONET data
- extend spectrally
- use simulation to scale back/forward in time
- add altitude information (modeling or CALIPSO)
- add information for associated CCN and IN
- some application: ... direct aero forcing in time

start with modeling

ensemble-median

advantages

 complete
 consistent
 (no data-gaps)

issues as good as assumptions

although usually 'tuned' to sat AOD observations

all relevant aerosol optical properties

- aerosol optical depth (AOD)
- single scattering albedo
- Angstrom (→ asymmetry factor)

annual aerosol optical depth at 550nm



improve with AERONET data

AERONET

- advantages
 high quality
- issues
 - local
 - global uneven

note, the climatology does NOT use satellite or in-situ data, they only constrain global modeling



the merging

– combine point data on regular grid (1°x1° lon/lat)

 extend spatial influence of identified sites with better regional representation

spatially spread valid grid ratios and combine

decaying over +/- 180 (Ion) and +/- 45 (Iat)

 – apply ratio field values (A) at site domains with distance decaying weights (B) → corr. factors



merging results



spectral extension

- use the Angstrom to stratify AOD into
 - AOD, f (fine sizes > 0.5μm radius, *Ap* =2.2-1.7)
 - AOD, c (coarse particles < 0.5μm rad., Ap=0)
- use SSA to establish dust size
 - dust is larger when fine SSA would fall below pre-scribed threshold
- coarse mode spectral dependence is defined
 composition and size was assigned
- fine mode spectral dependence only for solar
 - Ap, f define AOD, f spectral dependence
 - g (from Angstrom) and SSA drop in near-IR
spectral samples – annual avg maps

AOD

SSA

g



• VIS 0.55μm

• n-IR 1.0μm

・ IR 10μm



extension in time

- components that DO NOT change with time
 - coarse mode
 - pre-industrial fraction of fine mode
 - use info on what is 'anthropo' from modeling
- components that DO changes with time
 - stratospheric aerosol
 - anthropogenic fraction of the fine mode
 - back: ECHAM simulation with NIES emissions
 - forward: ECHAM runs with rcp scenarios

anthr.AOD,550nm - past

1865 to 2000



anthr.AOD,550nm - future

rcp_26 / rcp_45 / rcp_85



altitude

- current distributions
 - total aerosol
 - CALIPSO version 2
 - ECHAM
 - Separately for fine and coarse mode
 - ECHAM
- next ... CALIPSO version 3
 - use (non-sphere) (=dust) for coarse
 - use total-(non-sphere) for fine

total AOD (550nm) by layer

ECHAM 5 / HAM (model)

CALIPSO vers.2 (obs)



link to clouds

- aerosol provide CCN and IN to clouds
 - changes to aerosol concentration can modify cloud properties and the hydrological cycle
 - critical parameters are
 - aerosol concentration
 - aerosol composition
 - super-saturation
 - temperature

← from climatology← from modeling

CCN – need concentrations, not column data

aerosol concentrations

...are defined by the AOD vertical distribution and assumed log-normal size-distributions

- coarse-mode log-normal properties for seasalt or dust are prescribed
- fine-mode log-normal width is fixed (std.dev 1.7) and the mode radius is linked to the Angstrom of the fine-mode (which depends on low cloud cover)
 - larger aerosol radii at moister conditions
 - smaller aerosol radii at drier conditions

define CCN / IN

- all coarse mode particles are CCN ... plus
- a fraction of fine-mode particles are CCN
 - for fine-mode concentrations:
 - a critical cut-off size is determined as function
 - -super-saturation
 - kappa ('humidification') maps are based on ECHAM/HAM simulations for fine-mode AOD

» kappa weights: SS =1.0, SU =0.6, OC =0.1, DU/BC =0)

- only sizes larger that the cut-off are considered as fine-mode CCN
- coarse mode dust particles are IN

critical fine-mode radius

1km

3km

8km



CCN at 1.5km above ground & IN (ice nuclei) anthropogenic CCN concentrations 'dominate'



application – aerosol direct rad.forcing

- apply climatology in an off-line rad. transfer code for aerosol direct radiative forcing/effects
- focus on aerosol direct anthropogenic impact
 - current impact against natural impact(s)
 - changes in time (from 1875 to 2100)
- 'global' vs ('Mediterranean' or) 'E.Asia' impacts
- some results
 - natural impacts dominates
 - by 2100 global forcing at 1970 levels
 - by 2100 mediterranean forcing at 1910 levels
 - by 2100 E.Asia forcing at 1975 levels

radiative forcing

- the impact on the distribution of rad. energy ...
 - on the solar radiation (0.2 3.5 μ m)
 - on the terrestrial radiation (3.5 $100\mu m$)
- ... at the top of the atmosphere (ToA)
 - climate relevance: impact on the entire Earthatm-System: loss = cooling gain = warming
- ... at the surface
 - impact on surface temperature and processes
 (e.g. evap): loss = colder T gain = warmer T

aerosol radiative forcing

- impacts on the distribution of rad. energy ...
 - direct: by the presence of (added) aerosol
 - mainly a solar radiative (daytime) impact
 - indirect : (not covered) by modifying other atm properties which cause a forcing of their own
 changing cloud microphysics / cloud lifetime
- ... at the top of the atmosphere (ToA) compare to +2.8 W/m2 by anthrop greenh gases

aerosol direct forcing – ToA 70% of today's ToA aerosol forcing is natural

- aerosol on average "cools"
 - global avg: 1.6 W/m2
 - highly variable, even sign
 - anthrop. impact is much smaller: - 0.5 W/m2





aerosol direct forcing – SUrface 60% of today's surface aerosol forcing is natural

- aerosol always "cools"
 global average: 4.7W/m2
 - variable, source related
 - anthropogenic forcing is less than half: - 1.9 W/m2





natural vs. anthropogenic spatial pattern differs from anthropogenic



anthropogenic vs. natural - now (different x-axis scales)

anthropogenic monthly

natural monthly



past anthropogenic - ToA increased most rapidly in the 1950-1990 period



future anthropogenic - ToA will reduce by 2100 to about 65% of today's



anthropogenic forcing - summary

- today's forcing: 0.5W/m2 at ToA, -1.9 at surface
 - usually small compared to natural aero forcing
 - 30% (only 15% with Pinatubo) at ToA
 - 40% (only 30% with Pinatubo) at surface
- strongest increase in the 1950 1990
- should hit its maximum by 2020
- should return to 1970 levels by 2100
- the different warming scenarios are similar
 - 2100 ToA rcp 2.6 0.29 W/m2
 - 2100 ToA rcp 4.5 0.30 W/m2
 - 2100 ToA rcp 6.5 0.31 W/m2

global vs east-Asia

global

anthrop aerosol forcing:

- ToA average -0.47 W/m2
- ToA season -.36 to -.68
- surf average -1.9 W/m2
- maximum expected 2020
- predictions for 2100
 - 35% less than today
 - near 1970 levels
 - scenarios are similar

eastern Asia

anthrop aerosol forcing:

- ToA average -1.23 W/m2
- ToA season -.2 to -2.2
- surf average -6.2 W/m2
- maximum expected 2030
- predictions for 2100
 - 55% less than today
 - near 1975 levels
 - scenarios are similar

looking for trends ...







data access

ftp ftp-projects.zmaw.de

 cd aerocom/climatology/2010/yr2000 year 2000 · look at 'README_nc'

– ss-properties for the 14 solar-bands of ECHAM6

- g30_sol.nc
 g30_coa.nc
 g30_coa.nc
 g30_pre.nc
 g30_pre.nc
 g30_ant.nc
 total aerosol
 coarse mode aerosol
 (yr 2000)
 (yr 1850)
 (yr 1850)
 (yr 2000)
- ss-properties for the 16 IR-bands of ECHAM6
 - g30_fir.nc total (=coarse mode) aerosol
- CCN and IN fields
 - g_CCN_km20.nc for supersat. of 0.1, 0.2, 0.4 and 1% – anthrop. CCN is changing along with anthrop. AOD

data access

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anthropogenic change

- cd aerocom/climatology/2010/ant_historic
 - based on decadal change by ECHAM/HAM simulations for fine mode aerosol with NIES historic emissions
 - aeropt_kinne_550nm_fin_anthropAOD_yyyy.nc
 » simulations by Silvia Kloster, prep. by Declan O'Donnell
- cd aerocom/climatology/2010/ant_future
 - based on sulfate and carbon emission related regional changes to fine-mode AOD patterns with ECHAM/HAM for IPCC RCP 2.6, RCP 4.5 and RCP 8.5 scenarios till 2100
 - aeropt_kinne_550nm_fin_anthropAOD_rcp??_yyyy.nc
 » simulations by Kai Zhang, concepts by Hauke and Bjorn

data access

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altitude

- cd aerocom/climatology/2010/altitude
 - based on CALIPSO version2 multi-annual data
 - fine/coarse mode altitude differences via ECHAM/HAM
 - calipso_fc.nc
 - » CALIPSO data via Dave Winker (NASA_LaRC)
 - » ECHAM/HAM altitude data from Philip Stier
- cd aerocom/climatology/2010/ccn_historic
 - applying anthropogenic change to CCN concentrations
 - CCN_1km_yyyy.nc
 - CCN_3km_yyyy.nc
 - CCN_8km_yyyy.nc

at 1km above ground at 3km altitude at 8km altitude





direct forcing/effects global patterns

 clear ToA

 clear surface

 all-sky ToA

 all-sky surface

solar + IR

solar

anthrop. solar



direct ToA forcing – seasonality



← cooling

warming \rightarrow

direct forcing - sensitivity tests

- ToA all-sky forcing uncertainty
 - examine impact
 by changing
 individual prop.
 to aerosol or
 environment
 - +/- 0.3 W/m2



anthrop. direct forcing in time

- changing pattern
 - 1940: -.16 W/m²
 - 1945: -.17 W/m²
 - 1950: -.19 W/m²
 - 1955: -.21 W/m²
 - 1960: -.24 W/m²
 - 1965: -.27 W/m²
 - 1970: -.31 W/m²
 - 1975: **-**.35 W/m²
 - 1980: -.38 W/m²
 - 1985: -.41 W/m²
 - 1990: -.43 W/m²
 - 1995: -.44 W/m²



strongest increase

CCN – climatology in use

CCN /ccm3 at 1km 0.1% super-saturation



 liquid water path comp.
 stronger aerosol indirect effect with the climatology

by U.Lohmann ETHZ



environmental data



AOD / SSA / ASY at 550nm



AEROSOL, anthrop.

AOD / SSA / ASY


AEROSOL, fine

AOD by layer



AEROSOL, coarse

AOD by layer

