# Aerosols and Clouds as Forcings and Feedbacks in Global Climate Change

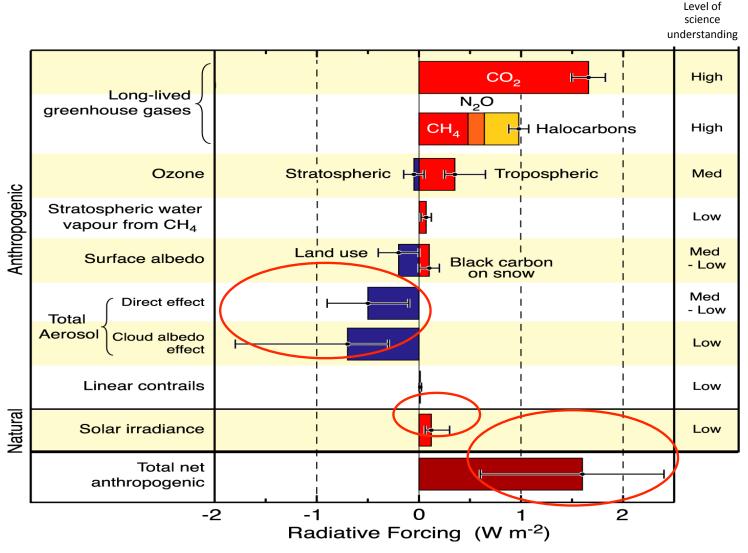
James Hansen

**12 September 2011** 

Université Pierre et Marie Curie

Paris, France

### Uncertainty in aerosol climate forcing is very large

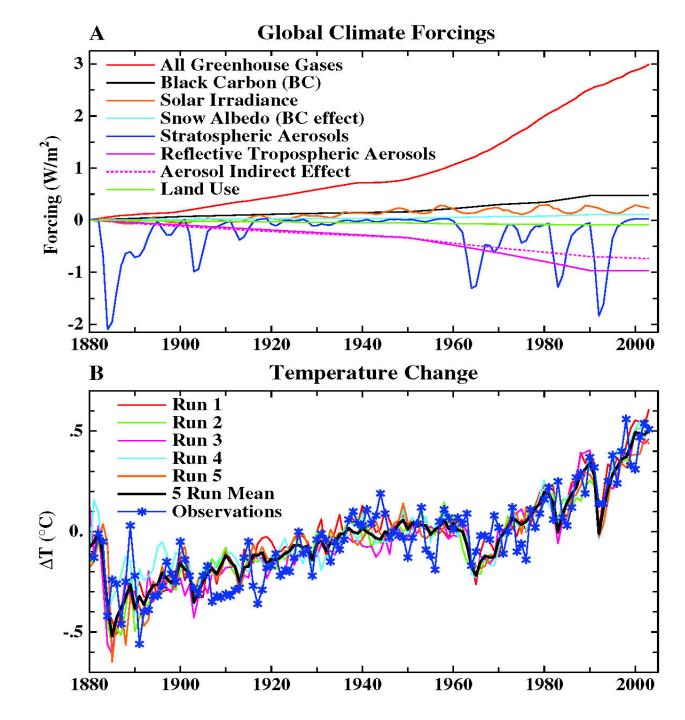


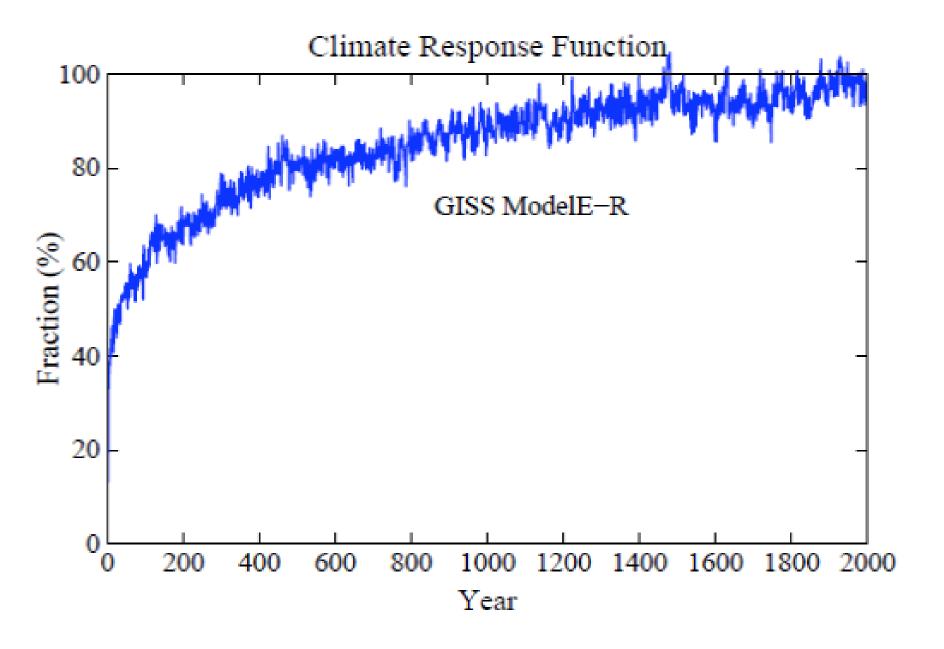
IPCC Summary for Policymakers, 2007

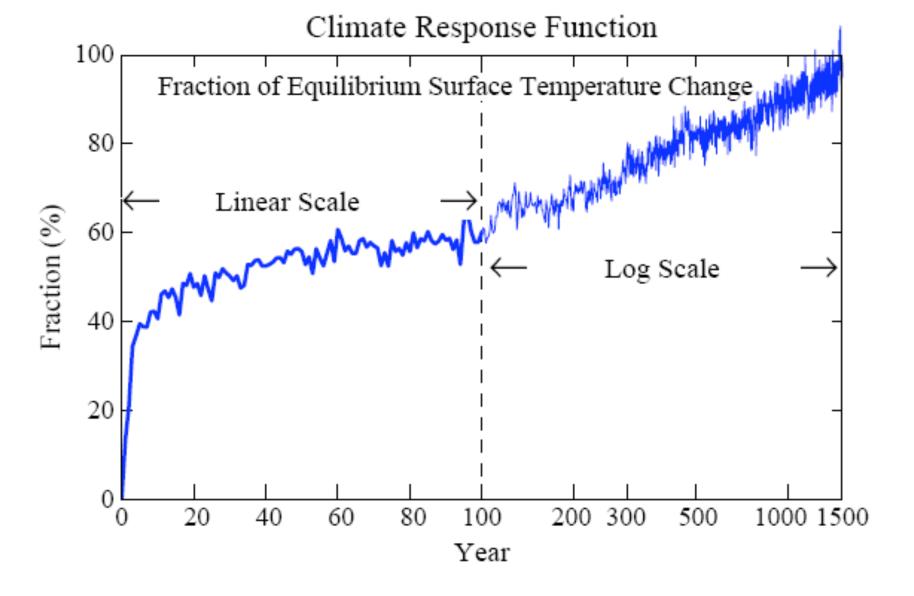
(A) Forcings used to drive climate simulations.

(B) Simulated and observed surface temperature change.

Source: Earth's energy imbalance: Confirmation and implications. Science **308**, 1431, 2005.

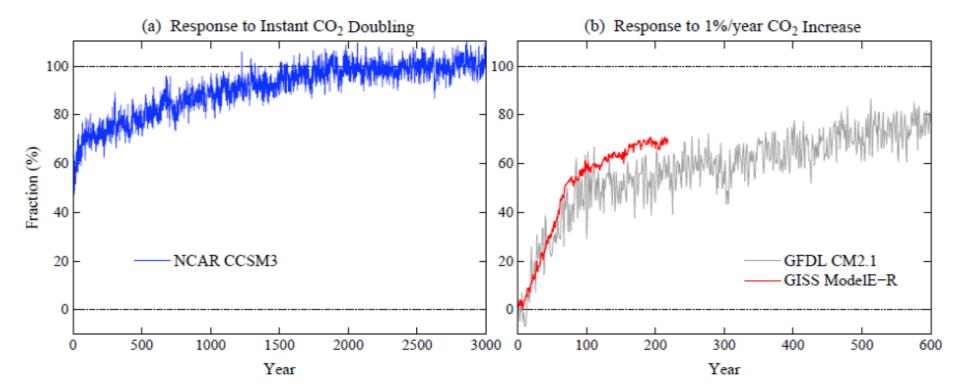






GISS climate model reaches only 60% of equilibrium response (to a given forcing) in a century. Other models have comparably long response times.

Source: Target Atmospheric CO<sub>2</sub>, Hansen et al., Open Atmos. Sci. J., 2, 217-231, 2008.



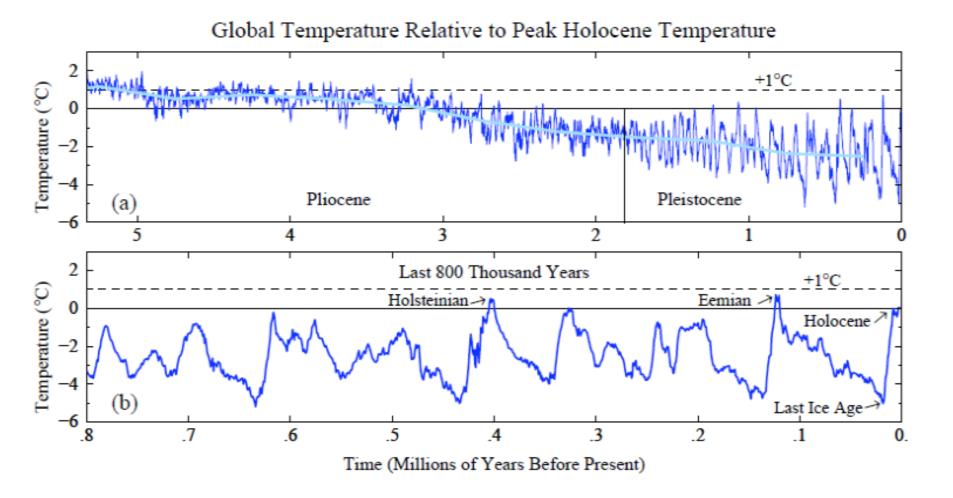
## **Feedbacks**

- 1. Respond to Climate, not Forcing
  - → response time ~ square of climate sensitivity
- 2. Slow Feedbacks

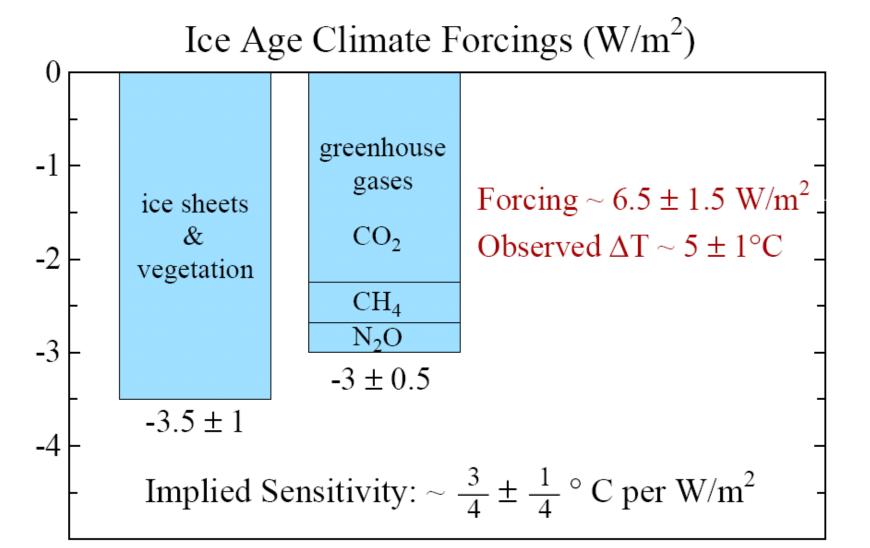
Further increase climate response time

3. Charney feedbacks vs. Fast feedbacks

Aerosols better treated as a fast feedback



**Figure 2.** Global temperature relative to peak Holocene temperature (Hansen and Sato, 2011).



Climate forcings during ice age 20 ky BP, relative to the present (pre-industrial) interglacial period.

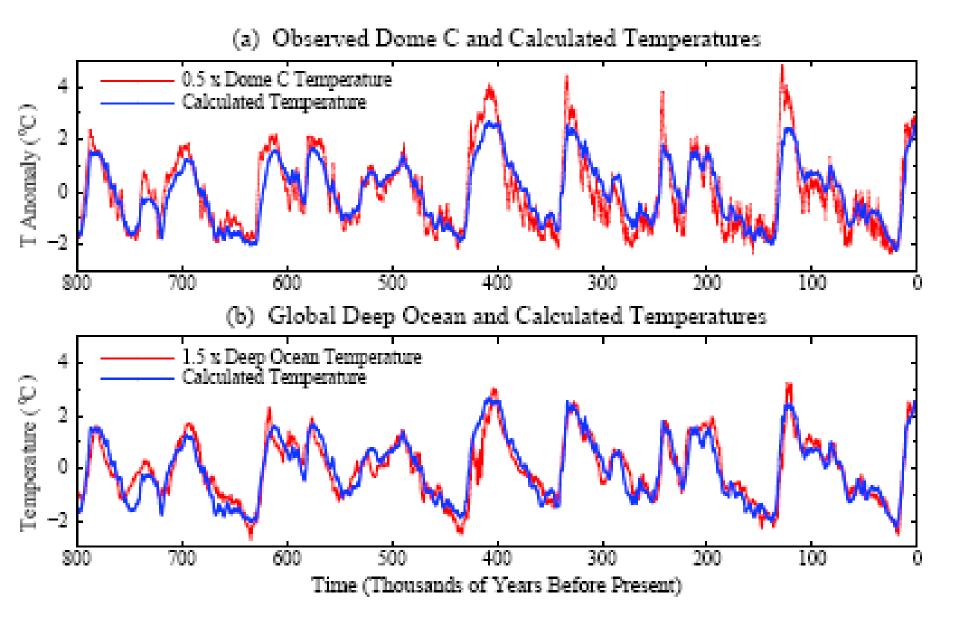


Figure 4

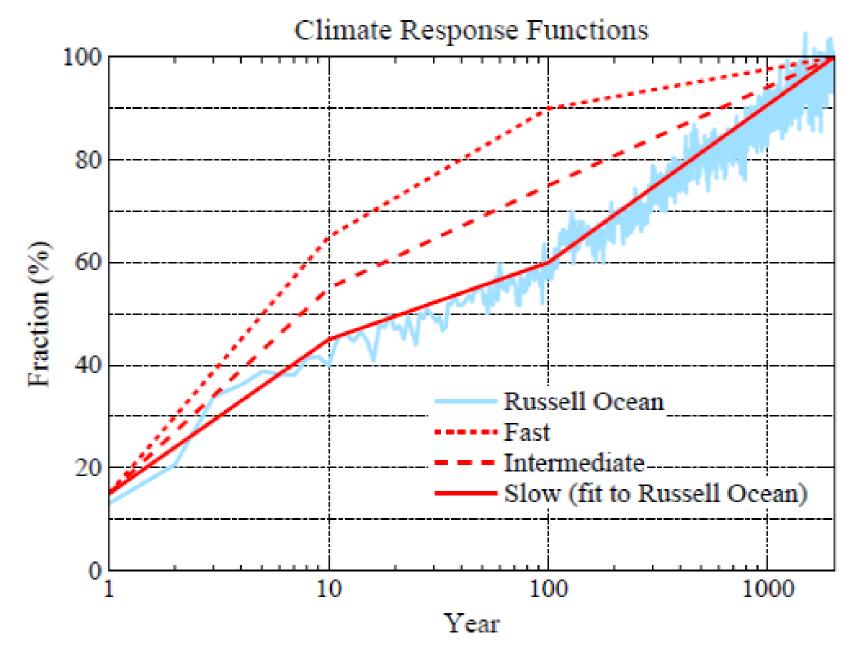


Figure 5

#### Greenhouse Gas, Aerosol & Net Climate Forcing

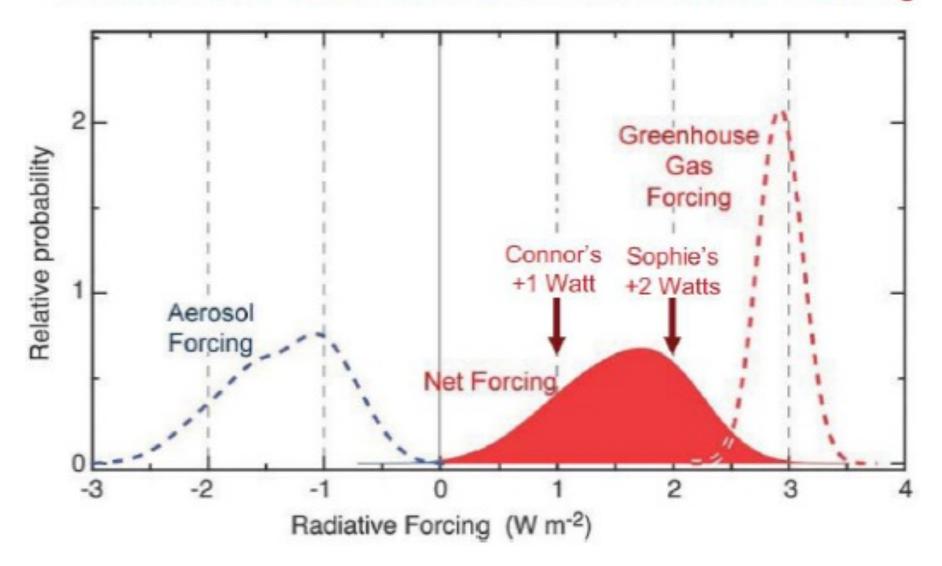
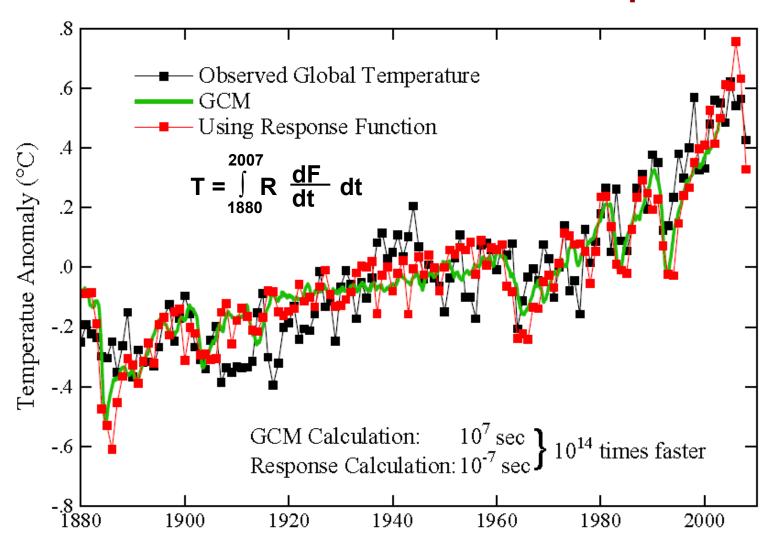


Figure 2

#### **Observed & Simulated Global Temperature**



Comparison of GCM results with simple integration of the Green's function [climate response function, R(t)] multiplied by forcing F(t).

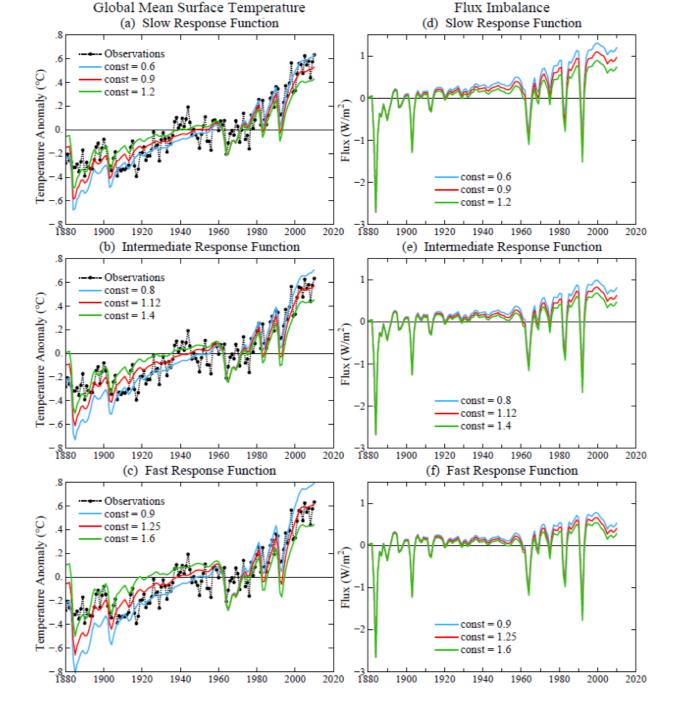
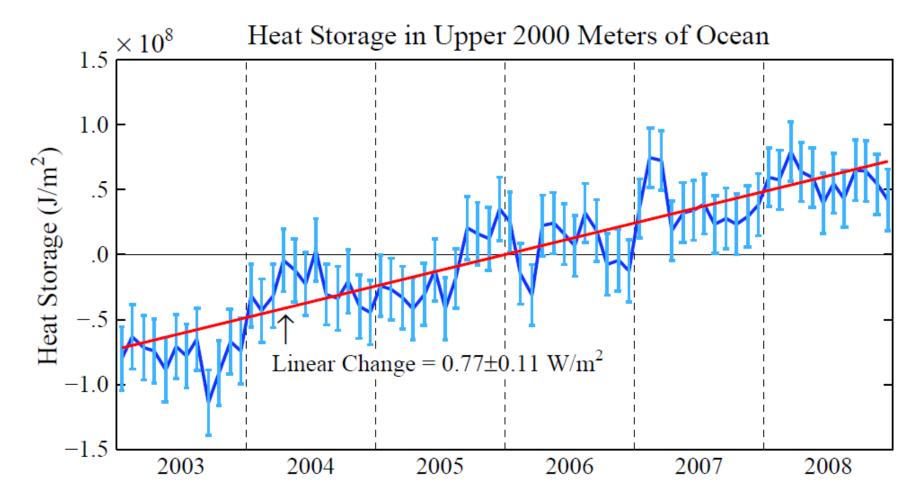


Figure 7



Heat storage in upper 2000 meters of ocean during 2003-2008 based on ARGO data. Knowledge of Earth's energy imbalance is improving rapidly as ARGO data lengthens. Data must be averaged over a decade because of El Nino/La Nina and solar variability. Energy imbalance is smoking gun for human-made increasing greenhouse effect.

Data source: von Schuckmann et al. J. Geophys. Res. 114, C09007, 2009, doi:10.1029/2008JC005237.

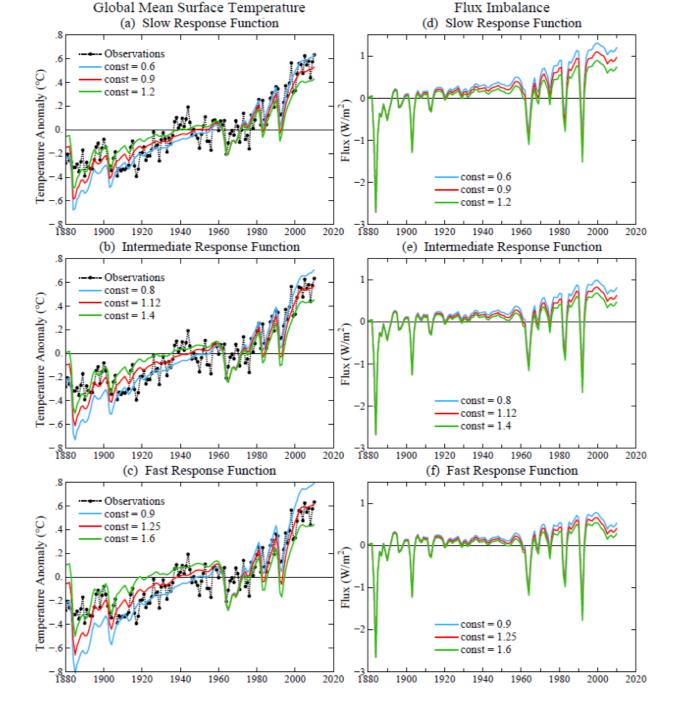
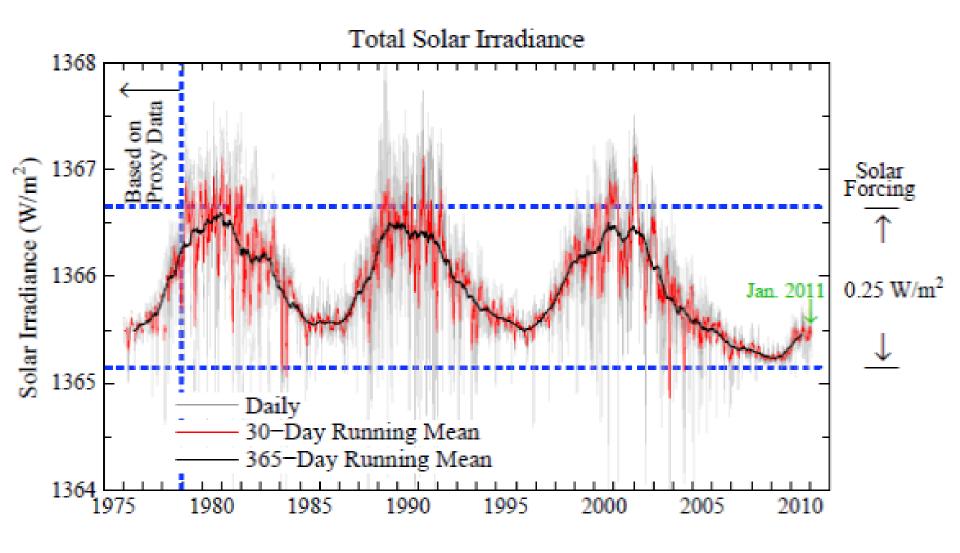


Figure 7



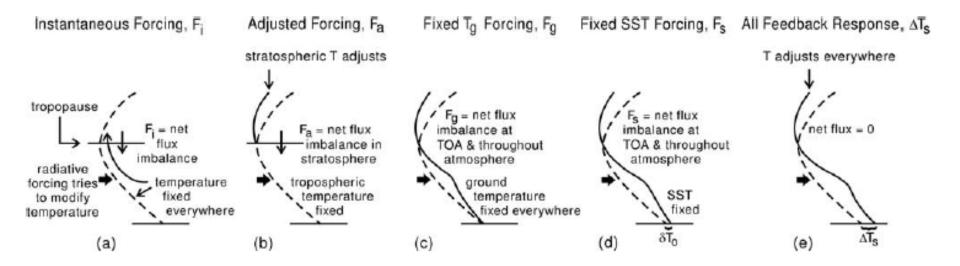


Figure 2. Cartoon comparing (a)  $F_i$ , instantaneous forcing, (b)  $F_a$ , adjusted forcing, which allows stratospheric temperature to adjust, (c)  $F_g$ , fixed  $T_g$  forcing, which allows atmospheric temperature to adjust, (d)  $F_s$ , fixed SST forcing, which allows atmospheric temperature and land temperature to adjust, and (e)  $\Delta T_s$ , global surface air temperature calculated by the climate model in response to the climate forcing agent.

# Forcings and Feedbacks

Forcing: Imposed Perturbation of Planet's Energy Balance

Feedback: Response of Climate System, which can Amplify or Diminish the Global Temperature Change Relative to Standard Radiative Change

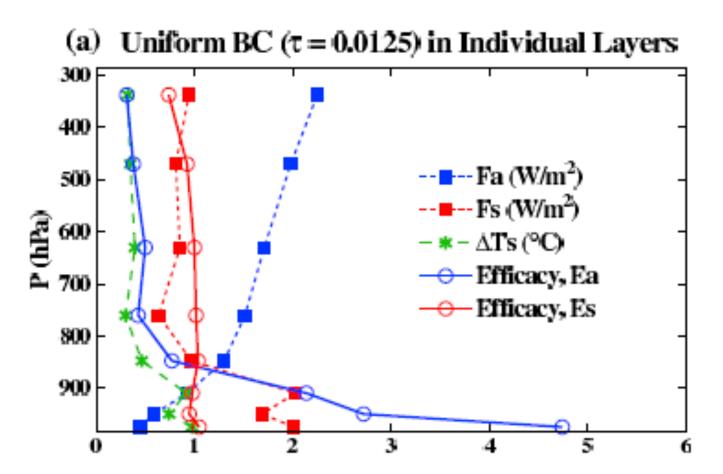


Figure 27. (a) Forcing, 81–120 year surface air temperature response, and efficacy for a globally uniform layer of BC aerosols in each of the eight lowest layers of the GISS model III. (b) Surface air temperature changes in years 81–

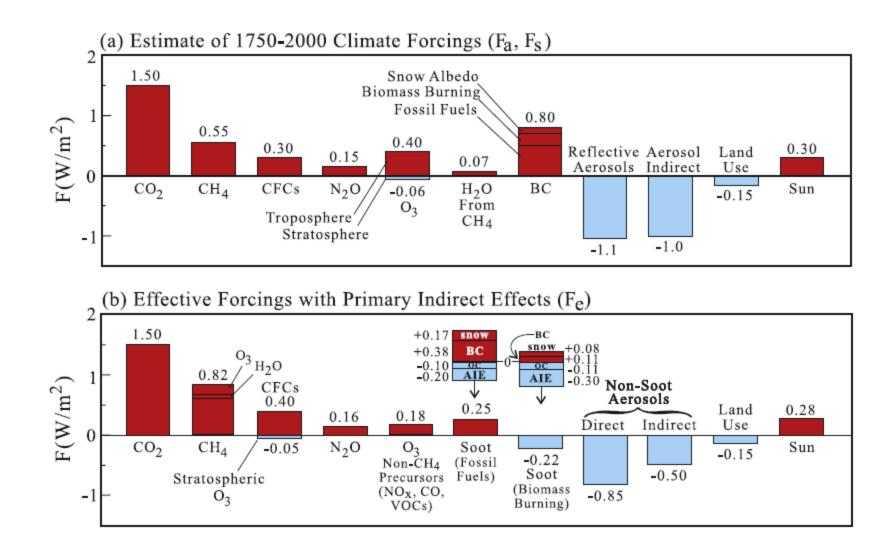


Figure 28. (a) A specific estimate of climate forcings for 1750–2000. (b) Same as Figure 28a, but with the effective forcing partially sorted by sources.

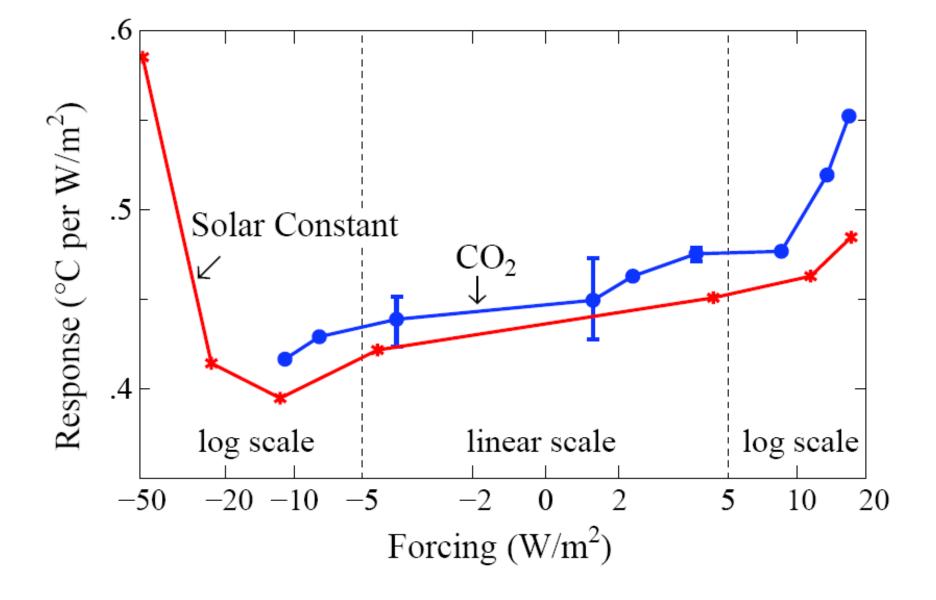


Figure S2 of "Target" paper.

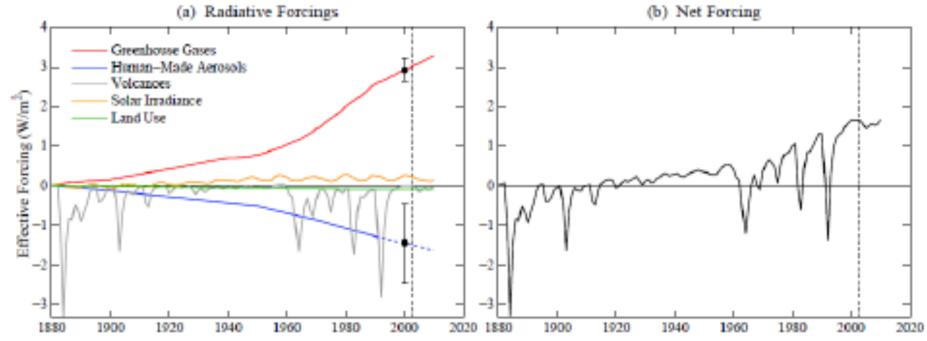


Fig. 1. Climate forcings employed in this paper. Forcings through 2003 (vertical line) are the same as used by Hansen et al. (2007), except the aerosol forcing after 1990 is approximated as -0.5 times the GHG forcing. Aerosol forcing includes all aerosol effects, including indirect effects on clouds and snow albedo. GHGs include O<sub>3</sub> and stratospheric H<sub>2</sub>O, in addition to well-mixed GHGs.

#### Planetary Energy Imbalance

